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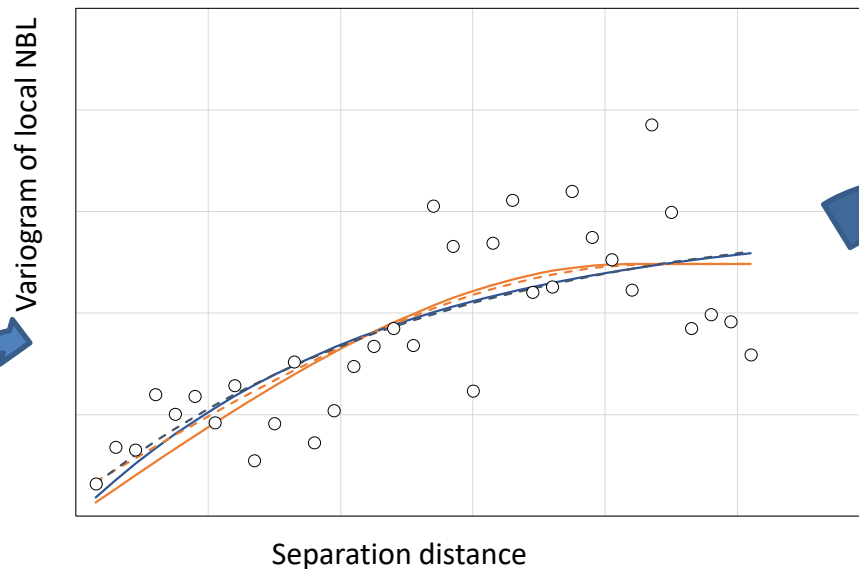
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Abstract: Quantification of the (spatially distributed) natural contributions to the chemical signature of groundwater resources is an emerging issue in the context of competitive groundwater uses as well as water regulation and management frameworks. Here, we illustrate a geostatistically-based approach for the characterization of spatially variable Natural Background Levels (NBLs) of target chemical species in large-scale groundwater bodies yielding evaluations of local probabilities of exceedance of a given threshold concentration. The approach is exemplified by considering three selected groundwater bodies and focusing on the evaluation of NBLs of ammonium and arsenic, as detected from extensive time series of concentrations collected at monitoring boreholes. Our study is motivated by the observation that reliance on a unique NBL value as representative of the natural geochemical signature of a reservoir can mask the occurrence of localized areas linked to diverse strengths of geogenic contributions to the groundwater status. We start from the application of the typical Pre-Selection (PS) methodology to the scale of each observation borehole to identify local estimates of NBL values. The latter are subsequently subject to geostatistical analysis to obtain estimates of their spatial distribution and the associated uncertainty. A multimodel framework is employed to interpret available data. The impact of alternative variogram models on the resulting spatial distributions of NBLs is assessed through probabilistic weights based on model identification criteria. Our findings highlight that assessing possible impacts of anthropogenic activities on groundwater environments with the aim of designing targeted solutions to restore a good groundwater quality status should consider a probabilistic description of the spatial distribution of NBLs. The latter is useful to provide enhanced information upon which one can then build decision-making protocols embedding the quantification of the associated uncertainty.

4 – Multi-model variogram assessment



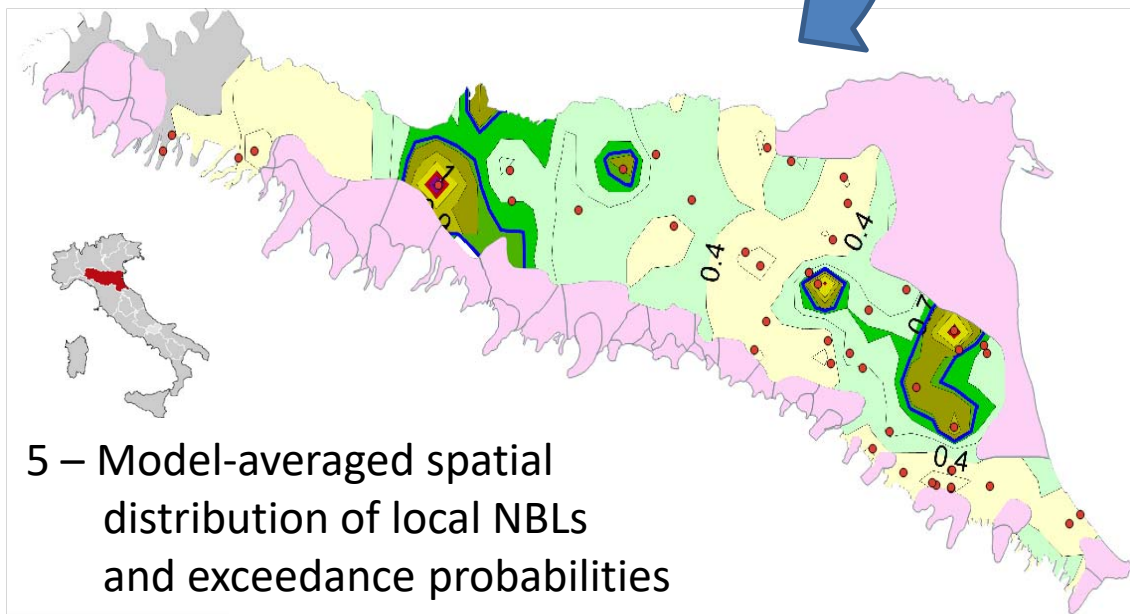
3 - Geospatial analysis



2 - Assessment of local NBL for each monitoring well



1 - Monitored concentrations



5 – Model-averaged spatial distribution of local NBLs and exceedance probabilities

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Geostatistical multimodel approach for the assessment of the spatial distribution of natural background concentrations in large-scale groundwater bodies

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Abstract

Quantification of the (spatially distributed) natural contributions to the chemical signature of groundwater resources is an emerging issue in the context of competitive groundwater uses as well as water regulation and management frameworks. Here, we illustrate a geostatistically-based approach for the characterization of spatially variable Natural Background Levels (NBLs) of target chemical species in large-scale groundwater bodies yielding evaluations of local probabilities of exceedance of a given threshold concentration. The approach is exemplified by considering three selected groundwater bodies and focusing on the evaluation of NBLs of ammonium and arsenic, as detected from extensive time series of concentrations collected at monitoring boreholes. Our study is motivated by the observation that reliance on a unique NBL value as representative of the natural geochemical signature of a reservoir can mask the occurrence of localized areas linked to diverse strengths of geogenic contributions to the groundwater status. We start from the application of the typical Pre-Selection (PS) methodology to the scale of each observation borehole to identify local estimates of NBL values. The latter are subsequently subject to geostatistical analysis to obtain estimates of their spatial distribution and the associated uncertainty. A multimodel framework is employed to interpret available data. The impact of alternative variogram models on the resulting spatial distributions of NBLs is assessed through probabilistic weights based on model identification criteria. Our findings highlight that assessing possible impacts of anthropogenic activities on groundwater environments with the aim of designing targeted solutions to restore a good groundwater quality status should consider a probabilistic description of the spatial distribution of NBLs. The latter is useful to provide enhanced information upon which one can then build decision-making protocols embedding the quantification of the associated uncertainty.

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1. Introduction

Modern society is characterized by an ever-increasing competitive use of groundwater resources, these being subject to many anthropogenic stresses (e.g., domestic use, irrigation and farming activities, industrial operations). To assist evaluation of the resilience of groundwater resources and the soil-water environment serving local communities, several studies are targeted to the analysis of water quality deterioration (e.g., Liu et al., 2017; Zlatanović et al., 2017; Heibati et al., 2017) or of water footprint characteristics (e.g., Qian et al., 2018). When dealing with the assessment of the qualitative status of a target groundwater body as a result of, e.g., in-place monitoring activities, it is not uncommon to evidence areas where detected chemical concentrations attain large values. In some instances, the latter can be directly or indirectly associated with the petrographical composition of the investigated aquifer (e.g., Hinsby and Condesso de Melo, 2006) or with site-specific characteristics such as the occurrence of organic matter (e.g., vegetal matter or peats) which can enhance release of chemical species to groundwater (e.g., Redman et al., 2002; Molinari et al., 2013 and references therein). These elements can in turn yield high natural levels of metals, such as Arsenic, even in crops (e.g., maize or rice) intended for human consumption (e.g., Kumarathilaka et al., 2018). Quantification of the actual (spatially distributed) natural (or geogenic) contributions to chemical concentration is an emerging issue causing increasingly pressing concerns in the context of competitive use of groundwater resources, water regulation and management frameworks at national and European levels, with implications in several industrial activities. Misclassifications of areas where sampled concentrations attain large values as a consequence of geogenic contributions yielding marked Natural Background Levels (NBLs) can have important socio-economic implications related to public health and risk assessment issues. Inaccurate risk assessment analyses can therefore yield an improper classification of the chemical status of an investigated aquifer which might lead to setting unrealistic remediation goals.

Characterization of the actual natural signature of groundwater bodies is a main theme of the EU Water Framework Directive (WFD 2000/60/EC, article 17). A key component required for the

67 reversal of identified marked and sustained upward trends of contaminants is the proper estimation
68 of NBLs of aquifer bodies. Main aspects related to the definition of NBLs are illustrated in article
69 2.5 of the GroundWater Daughter Directive (GWDD 2006/118/EC). The latter has been recently
70 amended by Directive 2014/80/EC stating that "the monitoring strategy and interpretation of the
71 data should take account of the fact that flow conditions and groundwater chemistry vary laterally
72 and vertically". With reference to these concepts, and prior to the enactment of Directive
73 2014/80/EC, Molinari et al. (2012) observe that (a) NBLs tend to increase with the average depth of
74 a water body and (b) whenever possible, NBLs should be estimated via robust experimental
75 characterization of the geochemical system and modeling studies performed, e.g., through state-of-
76 the-art multicomponent reactive transport approaches.

77 Statistical analysis of monitored data represents the typical approach employed for NBL
78 estimations (Edmunds et al., 2003; Wendland et al., 2005; Panno et al., 2006; Walter, 2008; Kim et
79 al., 2015). In this context, the EU research project BRIDGE (2007), Background cRiteria for the
80 IDentification of Groundwater thrEsholds, proposes a methodology termed as Pre-Selection (PS).
81 The latter is based on the identification of pristine groundwater samples across an available set of
82 sampled data, as representative of the natural population of the resident concentration. As a result of
83 this procedure, a unique (or bulk) NBL value is estimated and associated with the examined
84 subsurface reservoir, implying that all concentrations exceeding that level should be ascribed to
85 anthropogenic sources. The typical signature of a given chemical species in groundwater may be
86 defined through a range of concentrations rather than a single value (Reimann and Garrett, 2005;
87 Hinsby et al., 2008). This is related to the interaction and feedback between diverse natural,
88 atmospheric, geological, chemical and biological processes taking place in both the vadose and
89 saturated zone during groundwater infiltration and circulation (Edmunds et al., 2003; Wendland et
90 al., 2005; Panno et al., 2006; European Commission, 2009). These concepts are not completely
91 embedded in current regulatory frameworks which requires an estimate of only one threshold value,
92 considered as uniform across a given water body and against which anthropogenic contaminations

93 should be assessed (Reimann and Garrett, 2005). Otherwise, NBLs can attain markedly different
94 local values, for instance because of the occurrence of diverse petrographic provinces or redox
95 conditions within the same groundwater body, especially in large-scale reservoirs (with areal extent
96 of, e.g., thousands of square kilometers). Hence, the common practice of evaluating the chemical
97 status relying on a single NBL value cannot be considered as realistic and might lead to severe
98 over- or under-estimation of the typical natural signature.

99 Ducci et al. (2016) and Dalla Libera et al. (2017) recognize that the spatial distribution of
100 NBLs should reflect the heterogeneity of the investigated groundwater body. Critical assumptions
101 associated with these studies are (a) the reliance on a unique model employed to interpret
102 experimental variograms, (b) the incomplete quantification of the uncertainty of model parameters
103 and estimated concentrations, and (c) the lack of a direct estimation of local NBLs associated with
104 each monitoring well upon which exceedance probability maps can be conditioned. Yet, it is well
105 documented that estimated values and uncertainty analyses relying on a single (conceptual and/or
106 mathematical) model can lead to statistical bias or underestimation of the overall uncertainty linked
107 to the system behavior due to undersampling of the space of possible descriptive models. These
108 aspects can be seamlessly embedded within a Maximum Likelihood framework and subsequent
109 reliance on Model Quality criteria to consider uncertainty in the mathematical model depicting the
110 system as well as in its parameters (e.g., Carrera and Neuman, 1986; Ye et al., 2004, Bianchi Janetti
111 et al., 2012 and references therein; Gimeno et al., 2017).

112 Here, **our key objective is to** illustrate an approach for the estimation of local NBLs at the
113 borehole scale **through the application of** a geostatistically-based methodology to yield exceedance
114 probability maps. We **accomplish this objective by relying on** Maximum Likelihood and formal
115 model identification criteria to take into account uncertainty stemming from multiple and competing
116 models **which can be employed to interpret sample variograms of local values of NBLs**. To the best
117 of our knowledge, this approach stands as one of the first applications targeted at the evaluation of
118 NBL spatial maps by including **quantification of uncertainty associated with the variogram model**

119 employed to interpret the spatial distribution of local NBLs and probability exceedance
120 concentration maps, the latter being usually developed without the evaluation of local NBLs (e.g.,
121 Ungaro et al., 2008; Ayotte et al., 2006; Liu et al., 2004; Gaus et al., 2003).

122 **2. Materials and methods**

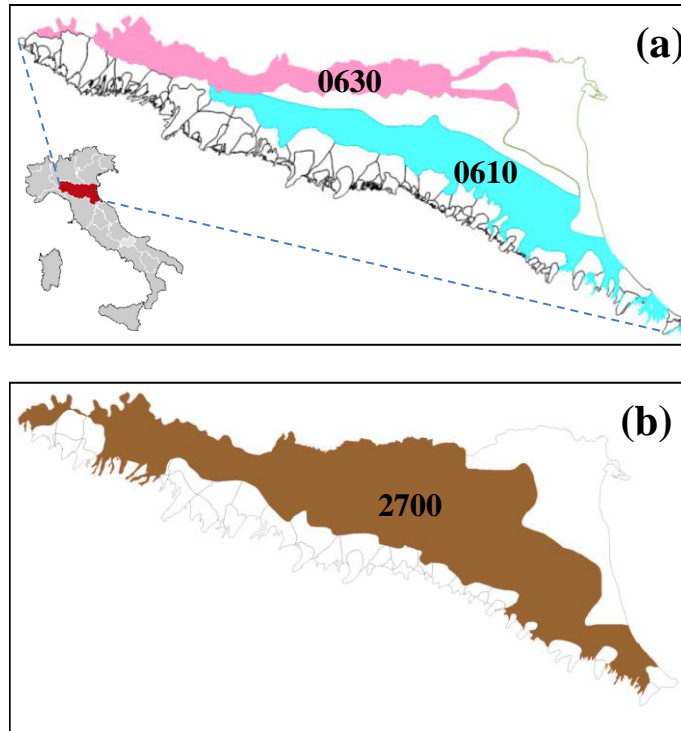
123 *2.1. Study areas*

124 Following the application of WFD 2000/60/EC, a total of 144 groundwater bodies have
125 been delimited within the Emilia-Romagna Region, Italy (Regione Emilia-Romagna, 2010). These
126 are part of the Po Basin fill, which is a syntectonic sedimentary wedge (Ricci Lucchi, 1984)
127 forming the infill of the Pliocene-Pleistocene fore-deep.

128 The information acquired from sedimentological and hydrogeological analyses has led to the
129 identification of three main hydrogeological complexes, i.e., Apennines alluvial fans, Apennine
130 alluvial plain and alluvial and deltaic Po plain. The system is characterized by a multilayered
131 confined or semiconfined configuration where gravel is gradually replaced by sand deposits in the
132 northern part of the plain, the thickness of fine deposits increasing towards the north portion of the
133 plain (Regione Emilia-Romagna, 2010; Farina et al., 2014). Additional information regarding
134 hydrogeological settings of the study region are available in Molinari et al. (2012) and Farina et al.
135 (2014). An upper confined portion and a lower confined portion have been distinguished within this
136 multilayered system. Our study is focused on the three largest groundwater bodies identified. Two
137 of these are located in the upper confined segment of the aquifer system of the Po Basin fill, the
138 remaining one representing a deep confined water body. Figure 1a depicts limits and planar extent
139 of the two upper confined groundwater bodies, respectively indicated as 0610 and 0630, while the
140 limits and planar extent of the deep confined water body, termed as 2700, are depicted in Figure 1b.

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 144 Fig. 1. Planar extent of the investigated water bodies in (a) the upper confined portion and (b) the
 145 lower confined portion of the aquifer system of the Po Basin fill; light blue = water body 0610, pink
 146 = water body 0630, brown = water body 2700.
 147

148 Table 1 lists the average depth and thickness as well as the planar area of the three groundwater
 149 bodies analyzed.

150

151

Table 1

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Extension and main characteristic length scales of the three groundwater bodies investigated.

Groundwater body	Average thickness (m)	Average depth (m)	Areal extent (km ²)
0610	130	75	2928
0630	110	65	1995
2700	180	200	6934

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154 Each of these groundwater bodies is subject to various levels of anthropogenic stresses
 155 because of diverse and competing uses of the subsurface, in terms of water consumption and
 156 withdrawals for agricultural and industrial purposes (Farina et al., 2014). Anthropogenic pollution
 157 tends to decrease going from superficial reservoirs to the deepest water body (i.e., 2700) which
 158 mostly receives diluted concentrations of chemical species from recharging areas located at some
 159 distance.

160 We consider these three systems because of their significant planar extent whose
161 representative scale is of the order of hundreds of kilometers and where the relevance of
162 considering regional-scale heterogeneous distributions of NBL values for the assessment of
163 groundwater quality is markedly evident.

164 2.2. Available dataset

165 We ground our analyses on the time series of concentrations collected at monitoring stations
166 associated with records of about 20 years of observations. These recordings (a) extend between
167 1987 and 2008 (albeit not continuously for some wells), (b) have been taken at a six-month interval,
168 and (c) constitute a unique data-base that we employ in the context of our investigation. The
169 chemical species considered in this study are ammonium (NH₄) and arsenic (As), which represent
170 critical elements for the achievement of a good chemical status for all three water bodies analyzed,
171 according to Italian Regulation (D. Lgs. 30/09, i.e., Decreto Legislativo n. 30, 16 March 2009) and
172 GWDD 2006/118/CE. As described in Molinari et al. (2012), to which we refer for further details,
173 all concentration data have been subjected to a preliminary exploratory statistical analysis that
174 identified ammonium and arsenic as critical species of concern.

175 Table 2 lists the number of monitoring stations and the total number of samples collected
176 within the 20-year long record of observations at locations included in the extensive network of
177 observation wells managed by the “Agenzia Regionale per la Prevenzione e l'Ambiente dell'Emilia-
178 Romagna” (ARPAE - Regional Agency for Environmental Protection, Emilia-Romagna).

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Table 2

Number of monitoring stations and total number of samples available

Groundwater body	monitoring stations	number of samples	
		As	NH ₄
0610	90	1968	2230
0630	75	1692	1917
2700	55	1201	1383

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2.3 Data analysis

2.3.1 NBL estimation

Within the framework of the EU research project BRIDGE (2007), the Pre-Selection (PS) methodology has been developed for the assessment of the overall geochemical signature of large-scale aquifer systems under data scarcity. The methodology relies on the statistical analysis of the information collected across a monitoring network and is based on the selection of samples that meet certain criteria and can be considered unaffected by anthropogenic influence. Typically adopted criteria for the exclusion of influenced samples are associated with the following conditions: (a) chloride concentrations > 1000 mg/L, as indicator of salinity; (b) nitrates (NO_3) concentrations > 10 mg/L, as indicator of human influence caused by, e.g., fertilizers; and (c) ammonium (NH_4) concentrations > 0.5 mg/L, as indicator of human impact under reducing conditions. Additional criteria, such as redox conditions, dissolved oxygen, sulfate concentration, can be considered for sample exclusion, as pointed out by Hinsby and Condesso de Melo (2006) and Hinsby et al. (2008) to maximize the possibility of grouping homogeneous data.

Samples exhibiting markers of anthropogenic contamination (e.g., nitrates or pesticides) larger than a given value have been removed from the original data bank and the residual set is used to estimate the median value for the remaining concentrations of the target chemical species at each monitoring well. The NBL value is then evaluated in terms of a selected percentile of the medians associated with each monitoring well within the investigated water body. The 90th, 95th, or 97.7th percentile are typically considered, depending on the degree of knowledge of the hydrogeochemical system. Wendland et al. (2005) propose to consider the 90th percentile of all of the calculated medians stemming from each monitoring well in the investigated reservoir as an estimate of NBL for the whole water body. Hereinafter, we refer to this quantity as $\text{NBL}_{90}^{\text{PS}}$.

208 As already stated, limitations inherent to the application of this procedure include: (a) all the
209 information associated with temporal variability of concentrations are shadowed, and (b) only one
210 NBL value is estimated for the whole aquifer body, without the possibility to assess any kind of
211 spatial variability across the system. With the aim of embedding within the analysis spatial and
212 temporal information linked to the scale of observation boreholes, we structure our study through
213 the following main steps:

- 214 1. perform sample selection for temporal records at each observation borehole following
215 typically adopted exclusion criteria, as **illustrated above and** indicated in the original
216 BRIDGE methodology;
- 217 2. evaluate a local NBL of the selected chemical species at each observation well as the 90th
218 percentile of concentration values retained at step 1;
- 219 3. perform a multimodel geostatistical analysis of the results from step 2, to (a) construct and
220 interpret empirical variograms of NBLs, (b) project local NBL values onto a computational
221 grid through Kriging and evaluate the associated variance, and (c) assess probabilities of
222 exceeding given threshold concentrations, considered as environmental performance metrics
223 characterizing the chemical status of the investigated system (see Section 2.3.2).

224 We emphasize that this approach directly imbues, as a result of step 2, the effects of the
225 monitored temporal variability of concentrations. As a result of step 3, spatial distributions of direct
226 local NBLs estimated for each observation well can be obtained, together with an appraisal of the
227 associated uncertainty, as reflected in the multimodel theoretical framework considered.

228 2.3.2 Spatial distribution of local NBLs

229 Following the approach described in Section 2.3.1 (step 3) we appraise the spatial
230 distribution of local NBLs within the target groundwater bodies through a geostatistical approach,
231 framed in the context of a Bayesian multimodel analysis. The study is performed according to the
232 following steps: (a) characterization of the spatial correlation structure of the variable by means of
233 experimental variograms; (b) selection of a set of alternative / competing theoretical variogram

234 models and estimation of their parameters (including their uncertainty) through Maximum
235 Likelihood (ML, see Section 2.3.3); (c) evaluation (through appropriate discrimination criteria and
236 posterior model weights, see Section 2.3.3) of the relative benefit associated with any of the models
237 considered to interpret available data; (d) projection of sample local NBLs onto a computational
238 grid via Kriging by relying on all of the calibrated models; (e) assessment of multimodel mean and
239 variance of local NBLs at each grid node (see Section 2.3.4); and (e) evaluation of the probability of
240 locally exceeding a given value, i.e., the uniform NBL value obtained through the PS procedure on
241 the regional scale.

242 A base 10 logarithmic transformation is applied to NBL concentration data to map these
243 onto the unbounded support comprising both positive and negative values. Omnidirectional
244 experimental variograms are assessed on the basis of the results of a preliminary variogram
245 analysis. Due to spatial arrangement of available sample points, Kriging estimates and variance are
246 calculated on a uniform grid, with spacing equal to 5 km. Geostatistical analyses has been
247 performed through the well-known and widely tested Stanford Geostatistical Modeling Software
248 (SGeMS; Remy et al., 2009).

249 We calibrate each of the models selected to interpret the evaluated sample variograms
250 through ML parameter estimation and apply model identification criteria to rank the tested models
251 in terms of posterior probabilistic weights. The latter are then used to weigh results associated with
252 each of the selected models and calculate multimodel mean and variance.

253 2.3.3 Maximum Likelihood (ML) parameter estimation and model quality criteria

254 Let N be the number of available observations of a model output Y collected in vector
255 $\mathbf{Y}^* = [Y_1^*, \dots, Y_N^*]$. Note that in our application these coincide with values of (log-transformed) NBLs
256 (see also Section 3). The covariance matrix of measurement errors, \mathbf{B}_Y , is here considered to be
257 diagonal with non-zero terms equal to the observation error variance, σ_i^2 (Carrera and Neuman,
258 1986). Denoting by $\hat{\mathbf{Y}} = [\hat{Y}_1, \dots, \hat{Y}_N]$ the vector of model predictions at locations where data are

259 available, the ML estimate $\hat{\mathbf{X}}$ of the vector of the M uncertain model parameters can be obtained
 260 by minimizing with respect to \mathbf{X} the negative log likelihood criterion:

$$261 \quad NLL = \sum_{i=1}^N \frac{J_i}{\sigma_i^2} + \ln |\mathbf{B}_Y| + N \ln(2\pi) \quad (1)$$

262 where $J_i = (Y_i^* - \hat{Y}_i)^2$. Criterion (1) includes the weighted least square criterion (Carrera and
 263 Neuman, 1986; Bianchi Janetti et al., 2012 and references therein). Here, minimization of (1) is
 264 achieved using the iterative Levenberg-Marquardt algorithm as embedded in the well documented
 265 computational framework PEST (Doherty, 2002).

266 Alternative (competing) models which can be used to interpret available system states can be
 267 ranked by various criteria (e.g., Neuman, 2003; Ye et al., 2004, 2008; Riva et al., 2011; Bianchi
 268 Janetti et al., 2012; Ciriello et al., 2015 and references therein), including:

$$269 \quad AIC = NLL + 2M \quad (2)$$

$$270 \quad AIC_c = NLL + 2M + \frac{2M(M+1)}{N-M-1} \quad (3)$$

$$271 \quad KIC = NLL + M \ln\left(\frac{N}{2\pi}\right) - \ln |\mathbf{Q}| \quad (4)$$

272 Here, \mathbf{Q} represents the Cramer-Rao lower-bound approximation for the covariance matrix of the
 273 parameter estimates, i.e., the inverse expected Fisher information matrix, which renders a
 274 quantitative appraisal of the quality of parameter estimates and of the information content carried by
 275 data about model parameters (see, e.g., Ye et al., 2008 for details). The Akaike information
 276 criterion, AIC, is due to Akaike (1974), AIC_c to Hurvich and Tsai (1989) and KIC to Kashyap
 277 (1982). The lowest value of a given model identification criterion indicates the most favored model
 278 (according to the criterion itself) at the expense of the remaining models. Note that *KIC* tends to (a)
 279 penalize models proportionally to the number of their parameters, through the quantity
 280 $M \cdot \ln(N/2\pi)$ (Ye et al., 2008; Hernandez et al., 2006; Riva et al., 2011) and (b) favor models with
 281 smaller expected information content per observation, when considering models with equal

282 parameter numbers, minimum NLL values and prior probability of parameters linked to such a
 283 minimum (Ye et al., 2008). In light of these observations, we base the analyses presented in this
 284 study on KIC (4).

285 The discrimination criteria (2)-(4) can also be considered to assign posterior probability
 286 weights quantifying uncertainty associated with each of the tested models. The posterior probability
 287 linked to model M_k ($k = 1, \dots, N_M$, N_M being the number of interpreting models considered) is
 288 evaluated as (Ye et al., 2008):

$$289 \quad p(M_k | \mathbf{Y}^*) = \frac{\exp\left(-\frac{1}{2} \Delta IC_k\right) p(M_k)}{\sum_{i=1}^{N_M} \exp\left(-\frac{1}{2} \Delta IC_i\right) p(M_i)} \quad (5)$$

290 Here, $\Delta IC_k = IC_k - IC_{min}$, IC_k being either AIC (2), AIC_c (3) or KIC (4) and $IC_{min} = \min\{IC_k\}$ its
 291 minimum value calculated across the range of models examined; $p(M_k)$ is the prior probability
 292 associated with each model. One can set $p(M_k) = 1/N_M$. In case no prior information is available, all
 293 models being then characterized by the same prior probability.

294 Grounding our study on model identification criteria and the ensuing posterior probabilities (5)
 295 enables one to rank the models analyzed through their posterior probabilities and discriminate
 296 among them in a relative sense.

297 2.3.4 Multimodel Mean and Variance

298 We consider a collection \mathbf{M} of K mutually exclusive variogram models, M_k , upon which
 299 lead statistics, such as mean and variance/covariance, of NBL values are computed through Kriging
 300 at the nodes of a selected computational grid covering a given aquifer body. The models are
 301 uncertain, each of them being assigned the same prior probability $p(M_k)$. Variogram model M_k is
 302 employed in a Kriging framework to yield the mean (expectation) $E(\mathbf{Y} | \mathbf{D}, M_k)$ and the covariance
 303 $Cov(\mathbf{Y} | \mathbf{D}, M_k)$ of a vector \mathbf{Y} of random (log-transformed) NBL values, conditional on the prior
 304 data vector \mathbf{D} . The entries of the latter are evaluated according to the procedure illustrated in

305 Section 2.3.1. Averaging across the moments provided by all K variogram models renders the
 306 following (Bayesian-averaged) lead moments (Draper, 1995; Hoeting et al., 1999):

$$307 \quad E(\mathbf{Y} | \mathbf{D}) = \sum_{k=1}^K E(\mathbf{Y} | \mathbf{D}, M_k) p(M_k | \mathbf{D}) \quad (6)$$

$$308 \quad \begin{aligned} Cov(\mathbf{Y} | \mathbf{D}) &= \sum_{k=1}^K Cov(\mathbf{Y} | \mathbf{D}, M_k) p(M_k | \mathbf{D}) \\ &+ \sum_{k=1}^K [E(\mathbf{Y} | \mathbf{D}, M_k) - E(\mathbf{Y} | \mathbf{D})] \\ &\cdot [E(\mathbf{Y} | \mathbf{D}, M_k) - E(\mathbf{Y} | \mathbf{D})]^T p(M_k | \mathbf{D}) \end{aligned} \quad (7)$$

309 T denoting transpose. The conditional covariance $Cov(\mathbf{Y} | \mathbf{D})$, resulting from Bayesian model
 310 averaging (BMA), is the sum of a within- and between-model contribution. Posterior model
 311 probabilities, $p(M_k | \mathbf{D})$, are calculated according to (5) and weigh the contribution of model M_k
 312 to BMA moments. For the purpose of our demonstration, we evaluate $E(\mathbf{Y} | \mathbf{D}, M_k)$ and
 313 $Cov(\mathbf{Y} | \mathbf{D}, M_k)$ through Kriging performed upon relying on variogram model M_k characterized
 314 through ML parameter estimates.

315 **3. Results and discussion**

316 Prior to the application of the PS procedure, the correct attribution of each monitoring
 317 station to a target groundwater body has been assessed on the basis of the technical characteristics
 318 of the boreholes (e.g., position of the filters) and the available hydrogeological information (e.g., the
 319 depth of permeable layers). Note that the application of the exclusion criteria described in Section
 320 2.3.1 has been performed by disregarding NH_4 because the collected sample cores provide evidence
 321 of natural occurrence of paleo-peats (Amorosi et al., 1996; Cremonini et al., 2008) consistent with
 322 documented increasing NH_4 concentrations with depth in the three reservoirs investigated (see also
 323 Molinari et al., 2012).

324 Exclusion of samples associated with anthropogenic influence (according to the criteria
 325 listed above) is followed by identification and removal of outliers from the remaining data set,

326 yielding the temporal record subsequently employed for local NBL estimation at each control point.
 327 Outliers are here identified as high concentration values that lie outside an interval centered around
 328 the sample mean of the data remaining after pre-selection and of width equal to three times the
 329 associated standard deviation. Excluding these records from the analyses is consistent with the
 330 observation that the bulk local environmental behavior is not significantly influenced by isolated (in
 331 time) and significantly large concentrations recorded, for instance, as a result of unusual processes
 332 taking place only in a particular year (and linked to the occurrence of, e.g., extreme rains that can
 333 cause large infiltration enhancing remobilization and diffusion of natural compounds), as opposed
 334 to a typical environmental baseline observed across the entire monitoring period.

335 Application of PS to the time series of each monitoring well is inevitably linked to a
 336 reduction of the number of records to be then subject to statistical analyses. We then ground our
 337 NBL calculations solely on monitoring wells which, following data selection, exhibit a time series
 338 with more than five records, characterizing an active monitoring period spanning at least 3 years
 339 (note that samples are collected on a six-month basis). Note that reliance on this kind of analysis is
 340 warranted for aquifers where a significant number of monitoring points is available with (a) a
 341 reliable and extensive temporal record of observations and (b) an appropriate degree of spatial
 342 coverage of the system. Table 3 lists the number of observation wells effectively employed for each
 343 water body after data selection.

344

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Table 3

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Number of observation wells available for spatial analysis of NBLs after data selection.

groundwater body	monitoring stations	monitoring stations
	Ammonium	Arsenic
0610	51	50
0630	62	60
2700	47	45

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When compared against Table 2, these data reveal that the upper confined water body 0610 suffers from a significant reduction (about 50%) of the total number of monitoring points due to

350 exclusion of data evidencing possible anthropogenic impact on the water body. Otherwise, only a
 351 limited reduction of observation wells is observed in the case of the other two water bodies (with a
 352 reduction of 20% and 16%, respectively for water body 0630 and 2700). With reference to
 353 anthropogenic pressures, these results are consistent with the observation that (a) the water body
 354 2700 is a deep reservoir and hence subject to more limited anthropogenic stresses than the other two
 355 bodies, which constitute upper confined reservoirs, (b) groundwater body 0630 is located in the
 356 proximity of the Po river, that can act as a source of recharge of the reservoir, thus reducing the
 357 impact of stresses caused, for instance, by pumping activities, and (c) reservoir 0610 is subject to a
 358 enhanced anthropogenic impact due to exploitation associated with intensive industrial/agricultural
 359 activities.

360 *3.1. Ammonium*

361 Values of local NBL_{90} obtained at each monitoring well through the methodology described
 362 in Section 2.3 display a high degree of spatial variability. These range from 0.828 to 20.835 mg/L
 363 (i.e., -0.08 to 1.319 in logarithmic scale) within groundwater body 0610, while ranging from 0.025
 364 to 14.406 mg/L (i.e., -1.602 to 1.159 in logarithmic scale) and from 0.025 to 30.362 mg/L (i.e., -
 365 1.602 to 1.482 in logarithmic scale), respectively within groundwater body 0630 and 2700. Values
 366 of NBL_{90}^{PS} estimated through the standard PS procedure at the regional scale (i.e., one value for each
 367 reservoir) for the three water bodies analyzed (Molinari et al., 2012) are listed in Table 4 together
 368 with the corresponding logarithmic-transformed values. We observe that all estimates exceed EU
 369 Drinking Water Standard for ammonium.

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Table 4
 Values for $NH_4 NBL_{90}$ estimated via the original PS procedure (NBL_{90}^{PS}). EU Drinking Water
 Standard for NH_4 concentration is 0.5 mg/L.

groundwater body	NBL_{90}^{PS} (mg/L)	$\text{Log}_{10} NBL_{90}^{PS}$
0610	4.6	0.66
0630	5.2	0.71
2700	12.0	1.08

374

375 For the purpose of our application, classical isotropic Spherical or Exponential variogram
 376 models, with or without nugget, are considered. Table 5 lists the results of variogram model
 377 calibration analysis. For each groundwater body, these include the estimated set of model
 378 parameters (i.e., nugget, range, and sill) based on minimization of NLL (1) and the associated
 379 posterior probabilities calculated through (4) and (5). As stated, the analyses have been performed
 380 by considering logarithmic transformation of values of NBL_{90} (termed as $NH_4Log_{10} NBL_{90}$).

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Table 5

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Estimated parameters of variogram models of $NH_4Log_{10} NBL_{90}$ values and associated posterior
 385 probability (p) based on KIC (4). Here, Sph = Spherical model, Exp = Exponential model; n =
 386 nugget, a = range (practical range for Exponential model), c = sill; (*) NBL_{90} values are given in
 387 (mg/L).

387

	Groundwater body 0610				Groundwater body 0630				Groundwater body 2700			
Model	n	a	c	p	n	a	c	p	n	a	c	p
	(*)	(km)	(*)	(%)	(*)	(km)	(*)	(%)	(*)	(km)	(*)	(%)
Sph	-	18.7	0.106	52	-	83.5	0.496	40	-	81.7	0.485	6
Exp	-	22.9	0.107	48	-	138.2	0.581	49	-	100.4	0.509	34
Sph	-	-	-	-	0.045	87.9	0.452	4.2	0.1	101.0	0.401	5
Exp	-	-	-	-	0.031	159.6	0.574	6.6	0.1	208.9	0.538	55

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The spatial variogram structure of $NH_4Log_{10} NBL_{90}$ within groundwater body 0610 has been
 interpreted by a Spherical and an Exponential model without nugget effect (introducing a nugget
 effect in the model calibration procedure is associated with near-zero estimated nugget and
 posterior weight and is therefore disregarded). Both variogram models exhibit similar values of
 range (practical range in the case of Exponential model) and sill, with nearly coinciding posterior
 probabilities.

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Four theoretical models, i.e., Spherical and Exponential models with and without nugget, are
 considered for groundwater bodies 0630 and 2700. The contribution of the nugget to the total
 variance is less than 10% in the case of groundwater body 0630. Otherwise, the estimated nugget

398 effect for groundwater body 2700 is equal to 15.7% and 20% of the total variance, respectively for
399 the Exponential and Spherical model with nugget, suggesting the occurrence of a significant degree
400 of variability between sample pairs at short distances.

401 Estimated values of correlation scale (i.e., the range) are lowest for water body 0610, a
402 finding that might be related to the spatial arrangement of sampling boreholes across the system or
403 to the observation that state variables, such as given quantiles of concentration time series observed
404 at multiple wells, are not necessarily characterized by a strong degree of spatial correlation, due to
405 the dynamics associated with the key processes driving chemical migration in the system.
406 Otherwise, a consistent spatial persistence of correlation structure of ammonium NBL_{90} is observed
407 in water bodies 0630 and 2700, where spatial variations of kriged estimates are then expected to be
408 quite smooth. The largest posterior probability values for water body 0630 are associated with
409 Spherical and Exponential models without nugget. On the other hand, the Exponential variogram
410 model (with or without nugget) is unambiguously favored in water body 2700.

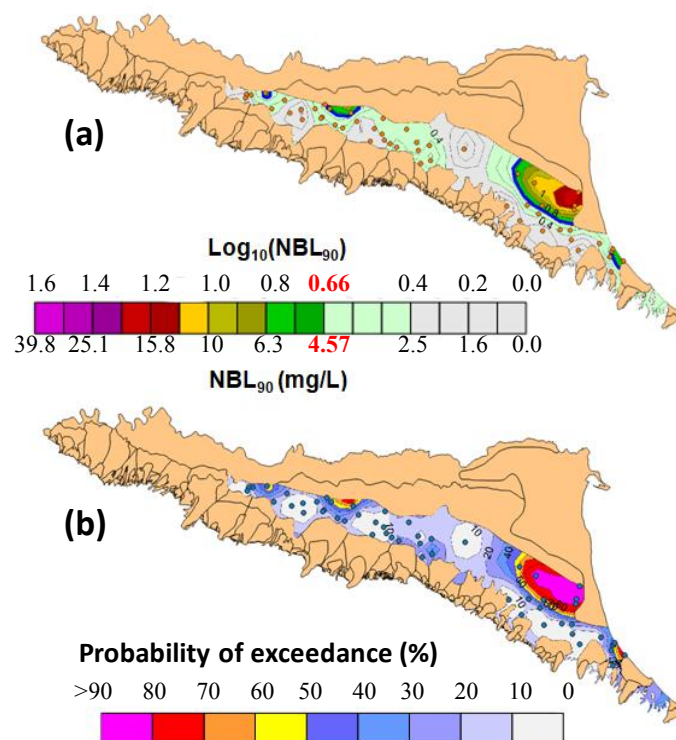
411 As described in Section 2.3, the calibrated variogram models are then employed to obtain
412 spatial distributions of Kriging estimates of log-transformed values of NBL_{90} and their associated
413 variance across each of the water bodies considered. For the purpose of this study we rely on
414 Ordinary Kriging, other flavors of Kriging being compatible with our approach. Kriging estimates
415 and variances obtained for each model are then weighted on the basis of posterior probability values
416 according to (6) and (7), to yield spatial distributions of estimated mean and variance of $\text{NH}_4\text{Log}_{10}$
417 NBL_{90} based on the complete set of tested models and fully including information on model
418 uncertainty.

419 As an example of the results obtained by considering individual variogram models, spatial
420 distributions of kriged values for local NBLs of ammonium in water body 2700 with the four
421 models listed in Table 5 are depicted in Appendix A. The spatial maps of kriged values of
422 $\text{NH}_4\text{Log}_{10}$ NBL_{90} for groundwater bodies 0610, 0630 and 2700 and resulting from our multimodel
423 analysis are respectively depicted in Figures 2, 3 and 4 together with the probability of exceeding

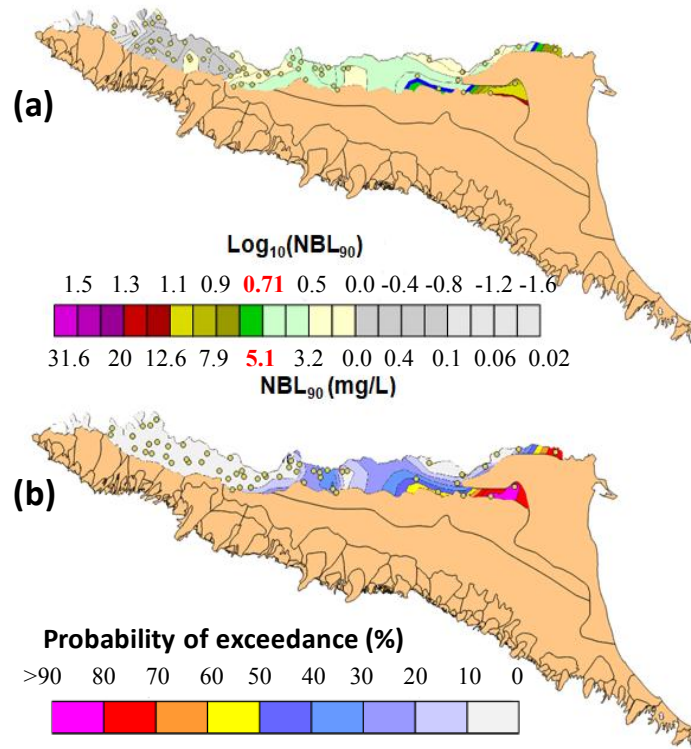
424 the uniform regional (log-transformed) NBL_{90}^{PS} value calculated via the original PS method (see
425 Table 4). For completeness, the corresponding spatial distributions of Kriging variance for the three
426 water bodies investigated are depicted in Appendix B. Exceedance probabilities are calculated by
427 assuming a Gaussian distribution of log-transformed NBL_{90} with local mean and variance rendered
428 by the multimodel analysis. Areas characterized by estimated mean values of $NH_4Log_{10}NBL_{90}$
429 larger than their uniformly distributed counterpart based on the original PS method are also
430 demarcated in Figures 2a, 3a, and 4a.

431 Spatial patterns of mean $Log_{10} NBL_{90}$ display a marked degree of variability across the
432 systems, suggesting that relying on a single (uniform) value can shadow the proper representation
433 of the actual NBL distribution within a given reservoir. For example, one can observe the
434 occurrence of areas where Kriging-based NBL_{90} exhibit values larger than NBL_{90}^{PS} . Otherwise, one
435 can also observe the occurrence of regions where estimated average NBLs are lower than NBL_{90}^{PS} ,
436 yet still larger than the EU drinking water standard set for ammonium. These observations reinforce
437 the idea that the use of a single NBL value as representative of the whole reservoir may lead to
438 misleading conclusions with reference to the natural behavior of the water body. For example, one
439 might argue that values larger than NBL_{90}^{PS} be associated with external causes (e.g., anthropogenic
440 activities) while they could be linked to specific and local hydrogeochemical natural processes.
441 Likewise, one should also consider that it is possible that concentration values in some areas be
442 lower than NBL_{90}^{PS} , yet larger than their regulatory-based counterpart (i.e., 0.5 mg/L in the case of
443 NH_4), as a result of natural processes rather than anthropogenically induced pollution phenomena.
444 These observations are also supported by the exceedance probability maps depicted in Figures 2b,
445 3b, and 4b. These suggest that there is non-negligible probability that NBL_{90}^{PS} be exceeded over a
446 large portion of the system, reinforcing the concept that the emergence of areas associated with a
447 geogenic origin of a target pollutant is masked when a single NBL_{90} value is taken as representative
448 of the whole reservoir. These sets of results represent an element upon which one could derive

449 information for the design of additional investigations. These could be directed, for example, to
 450 reduce uncertainty (as quantified, e.g., by large values of Kriging variance) in critical areas and/or
 451 to support the findings of the geostatistical analyses through the joint use of other types of
 452 information, including, e.g., (a) site-specific mineralogy, (b) the extent of petrographic provinces,
 453 or (c) the attainment of local natural hydrochemical equilibria linked to water-rock interactions with
 454 the aim of providing a complete picture of the natural signature of the system behavior. An analysis
 455 of this kind is outside the scope of the current study and will be the subject of future investigations.
 456

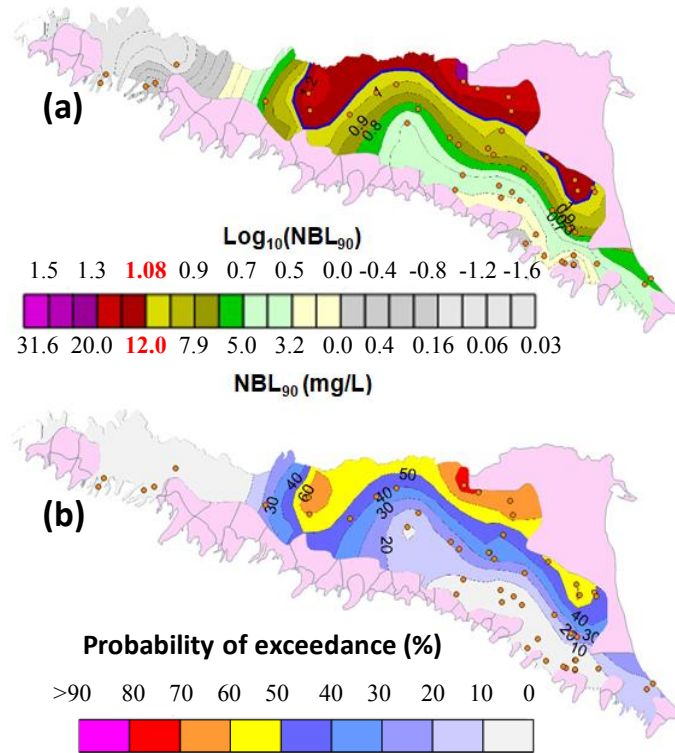


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 459 Fig. 2. Groundwater body 0610: spatial maps of (a) kriged $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}$ values and (b)
 460 probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the
 461 original PS method. The value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 0.66$ (red font on the color scale; corresponding
 462 to $\text{NBL}_{90}^{\text{PS}} = 4.6$ mg/L) is denoted by the thick blue isoline.
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Fig. 3. Groundwater body 0630: spatial maps of (a) kriged $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}$ values and (b) probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the original PS method. The value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 0.71$ (red font on the color scale; corresponding to $\text{NBL}_{90}^{\text{PS}} = 5.12 \text{ mg/L}$) is denoted by the thick blue isoline.



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Fig. 4. Groundwater body 2700: spatial maps of (a) kriged $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}$ values and (b) probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the original PS method. The value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 1.08$ (red font on the color scale; corresponding to $\text{NBL}_{90}^{\text{PS}} = 12.0 \text{ mg/L}$) is denoted by the thick blue isoline.

3.2 Arsenic

481 Values of local NBL_{90} obtained for arsenic (As) through the methodology described in
482 Section 2.3 display a considerable degree of spatial variability. These range from 1 to 120.4 $\mu\text{g/L}$
483 (from 0 to 2.08 in logarithmic scale) within groundwater body 0610, and from 1 to 49.8 $\mu\text{g/L}$ (from
484 0 to 1.7 in logarithmic scale) and from 0.5 to 70 $\mu\text{g/L}$ (from -0.3 to 1.84 in logarithmic scale),
485 respectively within reservoirs 0630 and 2700.

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Table 6
Values for As NBL_{90} estimated via the original PS procedure ($\text{NBL}_{90}^{\text{PS}}$). EU Drinking Water
Standard for As concentration is 10 $\mu\text{g/L}$.

groundwater body	$\text{NBL}_{90}^{\text{PS}}$ ($\mu\text{g/L}$)	$\text{Log}_{10} \text{NBL}_{90}^{\text{PS}}$
0610	33	1.52
0630	4	0.60
2700	6	0.77

490

491 Estimated values of NBL_{90}^{PS} for the three investigated water bodies (Molinari et al., 2012) are
 492 listed in Table 6 together with the corresponding log-transformed counterparts. We note that EU
 493 Drinking Water Standard for As (i.e., 10 $\mu\text{g/L}$) is exceeded only in the case of the 0610
 494 groundwater body.

495 Table 7 lists the results of the variogram model calibration analyses including, for each
 496 groundwater body, the estimated set of parameters based on minimization of NLL (1) and the
 497 associated posterior probabilities calculated through (4) and (5).

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Table 7

500 Estimated parameters of variogram models of $\text{AsLog}_{10} NBL_{90}$ values and associated posterior
 501 probability (p) based on KIC (4). Here, PN = Pure Nugget, Sph = Spherical model, Exp =
 502 Exponential model; n = nugget, a = range (practical range for Exponential model), c = sill; (*)
 503 NBL_{90} values are given in ($\mu\text{g/L}$).

Model	Groundwater body 0610				Groundwater body 0630				Groundwater body 2700			
	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)
PN	0.33	-	-	100	0.252	-	-	0.002	0.238	-	-	6.7
Sph	-	-	-	-	-	12.78	0.256	0.016	-	11.57	0.243	31.6
Exp	-	-	-	-	-	14.57	0.257	0.030	-	11.57	0.243	31.6
Sph	-	-	-	-	0.132	19.80	0.125	99.77	0.07	14.42	0.172	15.0
Exp	-	-	-	-	10^{-5}	14.58	0.257	0.177	0.07	14.42	0.172	15.0

504

505 We base our results on log-transformed values of local NBL_{90} (termed as $\text{AsLog}_{10} NBL_{90}$)
 506 estimated for each control borehole. The data observed for water body 0610 are characterized by a
 507 lack of spatial correlation structure, a pure nugget effect being the only model of choice with the
 508 ability to interpret the available data. This result is consistent with the low estimated values for the
 509 range of the variogram models employed to characterize the spatial correlation structure of
 510 ammonium NBL_{90} in the same water body (see Table 5). As observed in the case of ammonium,
 511 these results might be related to (a) the lack of sampling points separated by sufficiently small
 512 length scales or (b) the nature of the variable analyzed, which might or might not display significant
 513 spatial correlation structure. As a consequence, kriged local NBL values coincide with the mean of

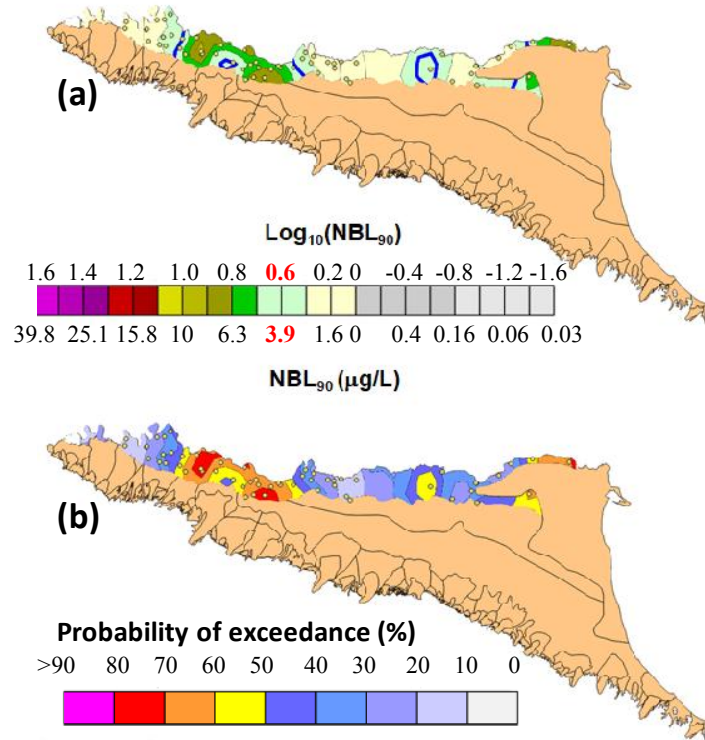
514 the available data, the associated variance being equal to the nugget. For these reasons, spatial maps
515 of $AsNBL_{90}$ for water body 0610 are not further analyzed.

516 The set of theoretical models employed to interpret the sample spatial correlation structure
517 of $AsNBL_{90}$ across groundwater bodies 0630 and 2700 comprises a pure nugget variogram model,
518 and spherical or exponential models with and without nugget. In the case of water body 0630,
519 model discrimination criteria assign the largest probabilistic weight (very close to 100%) to the
520 spherical model, which is also the model characterized by the largest contribution of the nugget to
521 the total variance (close to 50%). The other tested models are characterized by essentially negligible
522 weights. In the case of water body 2700, our results indicate that all models tested are associated
523 with a non-negligible probabilistic weight, the highest scores being equally assigned to the spherical
524 and exponential models without nugget. It is noted that addition of a nugget effect contributes to
525 about 28% of the total variance of the variogram models analyzed. Comparison of estimated
526 variogram model parameters obtained for arsenic and ammonium shows that variogram ranges for
527 As are significantly lower than those linked to NH_4 . This suggests that the spatial distribution of
528 values of $AsNBL_{90}$ may be driven by phenomena occurring at a more localized scale than in the
529 case of NH_4 .

530 We apply Ordinary Kriging to obtain Kriging estimates and variances of $AsNBL_{90}$ for each
531 variogram model. These are then weighted via posterior probability values according to (6) and (7),
532 to yield spatial distributions of estimated mean and variance of $AsLog_{10}NBL_{90}$ grounded on the
533 complete set of tested models. As an example of the results obtained by considering individual
534 variogram models, spatial distributions of kriged values for local NBLs of arsenic in water body
535 2700 with the four models listed in Table 5 are depicted in Appendix C.

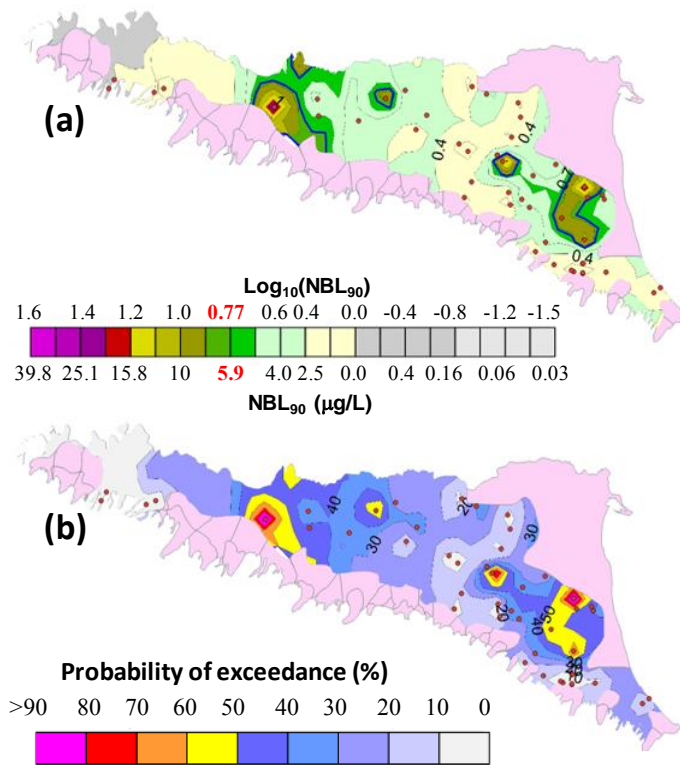
536 Kriged values of $AsLog_{10}NBL_{90}$ obtained from our multimodel analysis and associated with
537 groundwater bodies 0630 and 2700 are respectively depicted in Figures 5 and 6 together with the
538 probability of exceeding the corresponding uniform regional (log-transformed) NBL_{90}^{PS} value (see
539 Table 6). Exceedance probabilities are calculated following the procedure outlined in Section 3.1.

540 Areas characterized by estimated mean values of $AsLog_{10}NBL_{90}$ larger than their uniformly
 541 distributed counterpart based on the original PS method are also demarcated in Figures 5a and 6a.
 542 The corresponding spatial distributions of Kriging variance are depicted in Appendix D.
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 546 Fig. 5. Groundwater body 0630: spatial maps of (a) kriged $AsLog_{10}NBL_{90}$ values and (b)
 547 probability of exceedance of the uniform regional (log-transformed) NBL_{90}^{PS} value calculated via the
 548 original PS method. The value $AsLog_{10}NBL_{90}^{PS} = 0.60$ (red font on the color scale; corresponding to
 549 $NBL_{90}^{PS} = 4 \mu g/L$) is denoted by the thick blue isoline.

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555 Fig. 6. Groundwater body 2700: spatial maps of (a) kriged $\text{AsLog}_{10}\text{NBL}_{90}$ values and (b)
556 probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the
557 original PS method. The value $\text{AsLog}_{10}\text{NBL}_{90}^{\text{PS}} = 0.77 \mu\text{g/L}$ (red font on the color scale;
558 corresponding to $\text{NBL}_{90}^{\text{PS}} = 6 \mu\text{g/L}$) is denoted by the thick blue isoline.

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561 Similar to ammonium, we observe that considering a single NBL_{90} value as representative of
562 the whole water body can lead to underestimating/overestimating the actual natural signature which
563 can locally occur across the reservoir. These findings are markedly relevant for aquifers 0630 and
564 2700 where arsenic $\text{NBL}_{90}^{\text{PS}}$ values are lower than the EU drinking water standard, suggesting the
565 potential occurrence of localized areas characterized by local NBL values larger than the regulation
566 limit ($10 \mu\text{g/L}$). These observations are strengthened and quantified by the resulting exceedance
567 probabilities maps of arsenic (Fig. 5b and 6b). The latter could be consistent with the occurrence of
568 localized high natural arsenic content associated with the occurrence of vegetal matter, specific
569 solid fractions and redox conditions, as evidenced by Molinari et al. (2013, 2015).

569

4. Conclusions

570

Our work leads to the following major conclusions:

- 571 1. Relying on the estimate of a single (or bulk) NBL value in large-scale groundwater bodies
572 taken to represent the overall (median) behavior of the reservoir, tends to (a) mask the actual
573 distribution of local NBLs and (b) overestimate (or underestimate) natural background
574 concentrations within localized portions of the system. This might lead to **misleading**
575 conclusions within areas where values larger than the bulk NBL could be inappropriately
576 associated with external causes (e.g., anthropogenic activities) while they could **otherwise** be
577 linked to specific and localized natural processes. We show that probabilities of exceedance
578 of EU drinking water standard for hazardous species, such as arsenic, can vary quite
579 significantly across **a groundwater body. These elements evidence limitations associated**
580 **with current directives relying on a unique NBL value taken as representative of the bulk**
581 **behavior of an aquifer system. Such an approach can be especially critical in large-scale**
582 **groundwater systems (i.e., with planar extent of the order of thousands of square kilometers)**
583 **of the kind we analyze in this study. Considering our strategy, which is based on the**
584 **application of the typical Pre-Selection (PS) methodology to the scale of each observation**
585 **borehole where temporal records of monitored concentrations are available, yields local**
586 **estimates of NBL values. These can then be embedded in a geostatistical analysis to**
587 **characterize their spatial variability.**
- 588 2. By applying our modified PS approach, we estimate local scale NBLs, evaluated via the
589 information available at a given observation borehole. The interpretive multimodel approach
590 employed allows embedding explicitly uncertainties linked to the choice of the theoretical
591 model selected to describe the degree of spatial correlation (**as embedded in the variogram**
592 **model**) of such local scale NBL values. Relying on a single interpretive **variogram** model
593 can lead to an incomplete quantification of uncertainty, thus suggesting the need for
594 considering the contributions of a set of candidate models to underpin the incomplete
595 knowledge of the system behavior.

- 596 3. Spatial estimates of local NBLs and the ensuing uncertainty can serve as a basis to assist (i)
597 the demarcation of point and diffuse sources of potentially contaminating areas defining key
598 starting points and clean-up goals for remediation actions; and (ii) the management of
599 monitoring networks to optimize information content from areas associated with high
600 uncertainty or high probability of exceedance of environmental limits/standards.
- 601 4. Future European and national regulations related to NBLs assessment in large scale water
602 bodies should foresee the possibility of adopting spatial maps grounded on local NBLs
603 (based on temporal data sets). Otherwise, a unique value of NBL could be considered as
604 representative of the whole system for small scale reservoirs (planar extension lower than
605 hundreds of square kilometers) where limited amount of monitoring stations is typically
606 available and hampers a robust assessment of NBL spatial maps (with the associated
607 uncertainty).

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617 Data for the geostatistical analyses are available at the following link:
618 <https://data.mendeley.com/datasets/km6fgx227j/draft?a=3bdb20a5-adc7-429e-b1d0-3a12bd4202d8>.

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Appendix A. Supplementary data

776 Spatial distributions of kriged values for local NBLs of ammonium in water body 2700 with the
777 four models listed in Table 5 (the value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 1.08$ (red font on the color scale;
778 corresponding to $\text{NBL}_{90}^{\text{PS}} = 12.0$ mg/L) is denoted by the thick blue isoline.

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Appendix B. Supplementary data

781 Spatial distributions of Kriging variance obtained for ammonium for water bodies 0610, 0630 and
782 2700 as a result of the multimodel analysis.

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Appendix C. Supplementary data

785 Spatial distributions of kriged values for local NBLs of arsenic in water body 2700 with the four
786 models listed in Table 5. The value $\text{AsLog}_{10}\text{NBL}_{90}^{\text{PS}} = 0.77$ $\mu\text{g/L}$ (red font on the color scale;
787 corresponding to $\text{NBL}_{90}^{\text{PS}} = 6$ $\mu\text{g/L}$) is denoted by the thick blue isoline.

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Appendix D. Supplementary data

790 Spatial distributions of Kriging variance obtained for arsenic for water bodies 0630 and 2700 as a
791 result of the multimodel analysis.

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Geostatistical multimodel approach for the assessment of the spatial distribution of natural background concentrations in large-scale groundwater bodies

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Abstract

Quantification of the (spatially distributed) natural contributions to the chemical signature of groundwater resources is an emerging issue in the context of competitive groundwater uses as well as water regulation and management frameworks. Here, we illustrate a geostatistically-based approach for the characterization of spatially variable Natural Background Levels (NBLs) of target chemical species in large-scale groundwater bodies yielding evaluations of local probabilities of exceedance of a given threshold concentration. The approach is exemplified by considering three selected groundwater bodies and focusing on the evaluation of NBLs of ammonium and arsenic, as detected from extensive time series of concentrations collected at monitoring boreholes. Our study is motivated by the observation that reliance on a unique NBL value as representative of the natural geochemical signature of a reservoir can mask the occurrence of localized areas linked to diverse strengths of geogenic contributions to the groundwater status. We start from the application of the typical Pre-Selection (PS) methodology to the scale of each observation borehole to identify local estimates of NBL values. The latter are subsequently subject to geostatistical analysis to obtain estimates of their spatial distribution and the associated uncertainty. A multimodel framework is employed to interpret available data. The impact of alternative variogram models on the resulting spatial distributions of NBLs is assessed through probabilistic weights based on model identification criteria. Our findings highlight that assessing possible impacts of anthropogenic activities on groundwater environments with the aim of designing targeted solutions to restore a good groundwater quality status should consider a probabilistic description of the spatial distribution of NBLs. The latter is useful to provide enhanced information upon which one can then build decision-making protocols embedding the quantification of the associated uncertainty.

Keywords: Natural background levels; groundwater quality; chemical status; multimodel analyses; contaminated aquifers

1. Introduction

Modern society is characterized by an ever-increasing competitive use of groundwater resources, these being subject to many anthropogenic stresses (e.g., domestic use, irrigation and farming activities, industrial operations). To assist evaluation of the resilience of groundwater resources and the soil-water environment serving local communities, several studies are targeted to the analysis of water quality deterioration (e.g., Liu et al., 2017; Zlatanović et al., 2017; Heibati et al., 2017) or of water footprint characteristics (e.g., Qian et al., 2018). When dealing with the assessment of the qualitative status of a target groundwater body as a result of, e.g., in-place monitoring activities, it is not uncommon to evidence areas where detected chemical concentrations attain large values. In some instances, the latter can be directly or indirectly associated with the petrographical composition of the investigated aquifer (e.g., Hinsby and Condesso de Melo, 2006) or with site-specific characteristics such as the occurrence of organic matter (e.g., vegetal matter or peats) which can enhance release of chemical species to groundwater (e.g., Redman et al., 2002; Molinari et al., 2013 and references therein). These elements can in turn yield high natural levels of metals, such as Arsenic, even in crops (e.g., maize or rice) intended for human consumption (e.g., Kumarathilaka et al., 2018). Quantification of the actual (spatially distributed) natural (or geogenic) contributions to chemical concentration is an emerging issue causing increasingly pressing concerns in the context of competitive use of groundwater resources, water regulation and management frameworks at national and European levels, with implications in several industrial activities. Misclassifications of areas where sampled concentrations attain large values as a consequence of geogenic contributions yielding marked Natural Background Levels (NBLs) can have important socio-economic implications related to public health and risk assessment issues. Inaccurate risk assessment analyses can therefore yield an improper classification of the chemical status of an investigated aquifer which might lead to setting unrealistic remediation goals.

Characterization of the actual natural signature of groundwater bodies is a main theme of the EU Water Framework Directive (WFD 2000/60/EC, article 17). A key component required for the

67 reversal of identified marked and sustained upward trends of contaminants is the proper estimation
68 of NBLs of aquifer bodies. Main aspects related to the definition of NBLs are illustrated in article
69 2.5 of the GroundWater Daughter Directive (GWDD 2006/118/EC). The latter has been recently
70 amended by Directive 2014/80/EC stating that "the monitoring strategy and interpretation of the
71 data should take account of the fact that flow conditions and groundwater chemistry vary laterally
72 and vertically". With reference to these concepts, and prior to the enactment of Directive
73 2014/80/EC, Molinari et al. (2012) observe that (a) NBLs tend to increase with the average depth of
74 a water body and (b) whenever possible, NBLs should be estimated via robust experimental
75 characterization of the geochemical system and modeling studies performed, e.g., through state-of-
76 the-art multicomponent reactive transport approaches.

77 Statistical analysis of monitored data represents the typical approach employed for NBL
78 estimations (Edmunds et al., 2003; Wendland et al., 2005; Panno et al., 2006; Walter, 2008; Kim et
79 al., 2015). In this context, the EU research project BRIDGE (2007), Background cRiteria for the
80 IDentification of Groundwater thrEsholds, proposes a methodology termed as Pre-Selection (PS).
81 The latter is based on the identification of pristine groundwater samples across an available set of
82 sampled data, as representative of the natural population of the resident concentration. As a result of
83 this procedure, a unique (or bulk) NBL value is estimated and associated with the examined
84 subsurface reservoir, implying that all concentrations exceeding that level should be ascribed to
85 anthropogenic sources. The typical signature of a given chemical species in groundwater may be
86 defined through a range of concentrations rather than a single value (Reimann and Garrett, 2005;
87 Hinsby et al., 2008). This is related to the interaction and feedback between diverse natural,
88 atmospheric, geological, chemical and biological processes taking place in both the vadose and
89 saturated zone during groundwater infiltration and circulation (Edmunds et al., 2003; Wendland et
90 al., 2005; Panno et al., 2006; European Commission, 2009). These concepts are not completely
91 embedded in current regulatory frameworks which requires an estimate of only one threshold value,
92 considered as uniform across a given water body and against which anthropogenic contaminations

93 should be assessed (Reimann and Garrett, 2005). Otherwise, NBLs can attain markedly different
94 local values, for instance because of the occurrence of diverse petrographic provinces or redox
95 conditions within the same groundwater body, especially in large-scale reservoirs (with areal extent
96 of, e.g., thousands of square kilometers). Hence, the common practice of evaluating the chemical
97 status relying on a single NBL value cannot be considered as realistic and might lead to severe
98 over- or under-estimation of the typical natural signature.

99 Ducci et al. (2016) and Dalla Libera et al. (2017) recognize that the spatial distribution of
100 NBLs should reflect the heterogeneity of the investigated groundwater body. Critical assumptions
101 associated with these studies are (a) the reliance on a unique model employed to interpret
102 experimental variograms, (b) the incomplete quantification of the uncertainty of model parameters
103 and estimated concentrations, and (c) the lack of a direct estimation of local NBLs associated with
104 each monitoring well upon which exceedance probability maps can be conditioned. Yet, it is well
105 documented that estimated values and uncertainty analyses relying on a single (conceptual and/or
106 mathematical) model can lead to statistical bias or underestimation of the overall uncertainty linked
107 to the system behavior due to undersampling of the space of possible descriptive models. These
108 aspects can be seamlessly embedded within a Maximum Likelihood framework and subsequent
109 reliance on Model Quality criteria to consider uncertainty in the mathematical model depicting the
110 system as well as in its parameters (e.g., Carrera and Neuman, 1986; Ye et al., 2004, Bianchi Janetti
111 et al., 2012 and references therein; Gimeno et al., 2017).

112 Here, our key objective is to illustrate an approach for the estimation of local NBLs at the
113 borehole scale through the application of a geostatistically-based methodology to yield exceedance
114 probability maps. We accomplish this objective by relying on Maximum Likelihood and formal
115 model identification criteria to take into account uncertainty stemming from multiple and competing
116 models which can be employed to interpret sample variograms of local values of NBLs. To the best
117 of our knowledge, this approach stands as one of the first applications targeted at the evaluation of
118 NBL spatial maps by including quantification of uncertainty associated with the variogram model

119 employed to interpret the spatial distribution of local NBLs and probability exceedance
120 concentration maps, the latter being usually developed without the evaluation of local NBLs (e.g.,
121 Ungaro et al., 2008; Ayotte et al., 2006; Liu et al., 2004; Gaus et al., 2003).

122 **2. Materials and methods**

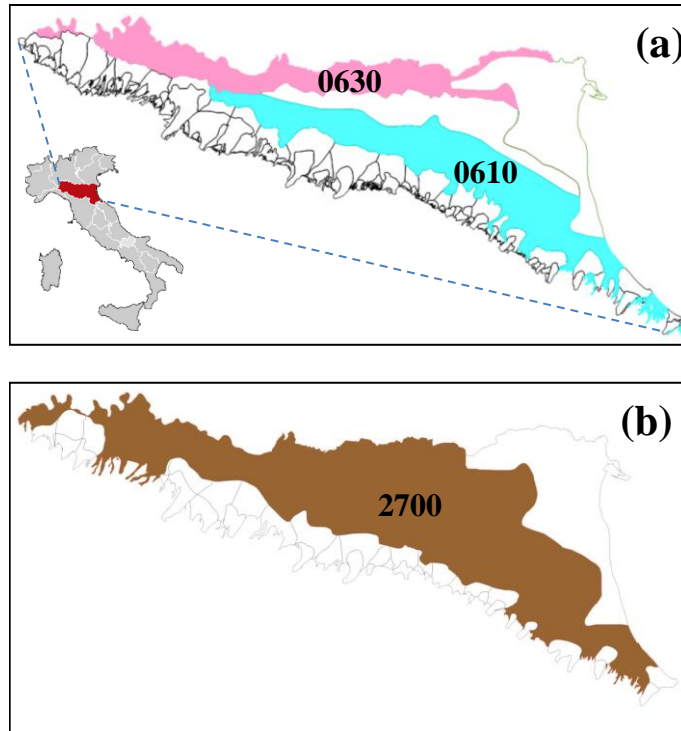
123 *2.1. Study areas*

124 Following the application of WFD 2000/60/EC, a total of 144 groundwater bodies have
125 been delimited within the Emilia-Romagna Region, Italy (Regione Emilia-Romagna, 2010). These
126 are part of the Po Basin fill, which is a syntectonic sedimentary wedge (Ricci Lucchi, 1984)
127 forming the infill of the Pliocene-Pleistocene fore-deep.

128 The information acquired from sedimentological and hydrogeological analyses has led to the
129 identification of three main hydrogeological complexes, i.e., Apennines alluvial fans, Apennine
130 alluvial plain and alluvial and deltaic Po plain. The system is characterized by a multilayered
131 confined or semiconfined configuration where gravel is gradually replaced by sand deposits in the
132 northern part of the plain, the thickness of fine deposits increasing towards the north portion of the
133 plain (Regione Emilia-Romagna, 2010; Farina et al., 2014). Additional information regarding
134 hydrogeological settings of the study region are available in Molinari et al. (2012) and Farina et al.
135 (2014). An upper confined portion and a lower confined portion have been distinguished within this
136 multilayered system. Our study is focused on the three largest groundwater bodies identified. Two
137 of these are located in the upper confined segment of the aquifer system of the Po Basin fill, the
138 remaining one representing a deep confined water body. Figure 1a depicts limits and planar extent
139 of the two upper confined groundwater bodies, respectively indicated as 0610 and 0630, while the
140 limits and planar extent of the deep confined water body, termed as 2700, are depicted in Figure 1b.

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Fig. 1. Planar extent of the investigated water bodies in (a) the upper confined portion and (b) the lower confined portion of the aquifer system of the Po Basin fill; light blue = water body 0610, pink = water body 0630, brown = water body 2700.

148 Table 1 lists the average depth and thickness as well as the planar area of the three groundwater
149 bodies analyzed.

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Table 1

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Extension and main characteristic length scales of the three groundwater bodies investigated.

Groundwater body	Average thickness (m)	Average depth (m)	Areal extent (km ²)
0610	130	75	2928
0630	110	65	1995
2700	180	200	6934

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154 Each of these groundwater bodies is subject to various levels of anthropogenic stresses
155 because of diverse and competing uses of the subsurface, in terms of water consumption and
156 withdrawals for agricultural and industrial purposes (Farina et al., 2014). Anthropogenic pollution
157 tends to decrease going from superficial reservoirs to the deepest water body (i.e., 2700) which
158 mostly receives diluted concentrations of chemical species from recharging areas located at some
159 distance.

160 We consider these three systems because of their significant planar extent whose
161 representative scale is of the order of hundreds of kilometers and where the relevance of
162 considering regional-scale heterogeneous distributions of NBL values for the assessment of
163 groundwater quality is markedly evident.

164 2.2. Available dataset

165 We ground our analyses on the time series of concentrations collected at monitoring stations
166 associated with records of about 20 years of observations. These recordings (a) extend between
167 1987 and 2008 (albeit not continuously for some wells), (b) have been taken at a six-month interval,
168 and (c) constitute a unique data-base that we employ in the context of our investigation. The
169 chemical species considered in this study are ammonium (NH₄) and arsenic (As), which represent
170 critical elements for the achievement of a good chemical status for all three water bodies analyzed,
171 according to Italian Regulation (D. Lgs. 30/09, i.e., Decreto Legislativo n. 30, 16 March 2009) and
172 GWDD 2006/118/CE. As described in Molinari et al. (2012), to which we refer for further details,
173 all concentration data have been subjected to a preliminary exploratory statistical analysis that
174 identified ammonium and arsenic as critical species of concern.

175 Table 2 lists the number of monitoring stations and the total number of samples collected
176 within the 20-year long record of observations at locations included in the extensive network of
177 observation wells managed by the “Agenzia Regionale per la Prevenzione e l'Ambiente dell'Emilia-
178 Romagna” (ARPAE - Regional Agency for Environmental Protection, Emilia-Romagna).

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Table 2

Number of monitoring stations and total number of samples available

Groundwater body	monitoring stations	number of samples	
		As	NH ₄
0610	90	1968	2230
0630	75	1692	1917
2700	55	1201	1383

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2.3 Data analysis

2.3.1 NBL estimation

Within the framework of the EU research project BRIDGE (2007), the Pre-Selection (PS) methodology has been developed for the assessment of the overall geochemical signature of large-scale aquifer systems under data scarcity. The methodology relies on the statistical analysis of the information collected across a monitoring network and is based on the selection of samples that meet certain criteria and can be considered unaffected by anthropogenic influence. Typically adopted criteria for the exclusion of influenced samples are associated with the following conditions: (a) chloride concentrations > 1000 mg/L, as indicator of salinity; (b) nitrates (NO_3) concentrations > 10 mg/L, as indicator of human influence caused by, e.g., fertilizers; and (c) ammonium (NH_4) concentrations > 0.5 mg/L, as indicator of human impact under reducing conditions. Additional criteria, such as redox conditions, dissolved oxygen, sulfate concentration, can be considered for sample exclusion, as pointed out by Hinsby and Condesso de Melo (2006) and Hinsby et al. (2008) to maximize the possibility of grouping homogeneous data.

Samples exhibiting markers of anthropogenic contamination (e.g., nitrates or pesticides) larger than a given value have been removed from the original data bank and the residual set is used to estimate the median value for the remaining concentrations of the target chemical species at each monitoring well. The NBL value is then evaluated in terms of a selected percentile of the medians associated with each monitoring well within the investigated water body. The 90th, 95th, or 97.7th percentile are typically considered, depending on the degree of knowledge of the hydrogeochemical system. Wendland et al. (2005) propose to consider the 90th percentile of all of the calculated medians stemming from each monitoring well in the investigated reservoir as an estimate of NBL for the whole water body. Hereinafter, we refer to this quantity as $\text{NBL}_{90}^{\text{PS}}$.

208 As already stated, limitations inherent to the application of this procedure include: (a) all the
209 information associated with temporal variability of concentrations are shadowed, and (b) only one
210 NBL value is estimated for the whole aquifer body, without the possibility to assess any kind of
211 spatial variability across the system. With the aim of embedding within the analysis spatial and
212 temporal information linked to the scale of observation boreholes, we structure our study through
213 the following main steps:

- 214 1. perform sample selection for temporal records at each observation borehole following
215 typically adopted exclusion criteria, as illustrated above and indicated in the original
216 BRIDGE methodology;
- 217 2. evaluate a local NBL of the selected chemical species at each observation well as the 90th
218 percentile of concentration values retained at step 1;
- 219 3. perform a multimodel geostatistical analysis of the results from step 2, to (a) construct and
220 interpret empirical variograms of NBLs, (b) project local NBL values onto a computational
221 grid through Kriging and evaluate the associated variance, and (c) assess probabilities of
222 exceeding given threshold concentrations, considered as environmental performance metrics
223 characterizing the chemical status of the investigated system (see Section 2.3.2).

224 We emphasize that this approach directly imbues, as a result of step 2, the effects of the
225 monitored temporal variability of concentrations. As a result of step 3, spatial distributions of direct
226 local NBLs estimated for each observation well can be obtained, together with an appraisal of the
227 associated uncertainty, as reflected in the multimodel theoretical framework considered.

228 2.3.2 Spatial distribution of local NBLs

229 Following the approach described in Section 2.3.1 (step 3) we appraise the spatial
230 distribution of local NBLs within the target groundwater bodies through a geostatistical approach,
231 framed in the context of a Bayesian multimodel analysis. The study is performed according to the
232 following steps: (a) characterization of the spatial correlation structure of the variable by means of
233 experimental variograms; (b) selection of a set of alternative / competing theoretical variogram

234 models and estimation of their parameters (including their uncertainty) through Maximum
235 Likelihood (ML, see Section 2.3.3); (c) evaluation (through appropriate discrimination criteria and
236 posterior model weights, see Section 2.3.3) of the relative benefit associated with any of the models
237 considered to interpret available data; (d) projection of sample local NBLs onto a computational
238 grid via Kriging by relying on all of the calibrated models; (e) assessment of multimodel mean and
239 variance of local NBLs at each grid node (see Section 2.3.4); and (e) evaluation of the probability of
240 locally exceeding a given value, i.e., the uniform NBL value obtained through the PS procedure on
241 the regional scale.

242 A base 10 logarithmic transformation is applied to NBL concentration data to map these
243 onto the unbounded support comprising both positive and negative values. Omnidirectional
244 experimental variograms are assessed on the basis of the results of a preliminary variogram
245 analysis. Due to spatial arrangement of available sample points, Kriging estimates and variance are
246 calculated on a uniform grid, with spacing equal to 5 km. Geostatistical analyses has been
247 performed through the well-known and widely tested Stanford Geostatistical Modeling Software
248 (SGeMS; Remy et al., 2009).

249 We calibrate each of the models selected to interpret the evaluated sample variograms
250 through ML parameter estimation and apply model identification criteria to rank the tested models
251 in terms of posterior probabilistic weights. The latter are then used to weigh results associated with
252 each of the selected models and calculate multimodel mean and variance.

253 2.3.3 Maximum Likelihood (ML) parameter estimation and model quality criteria

254 Let N be the number of available observations of a model output Y collected in vector
255 $\mathbf{Y}^* = [Y_1^*, \dots, Y_N^*]$. Note that in our application these coincide with values of (log-transformed) NBLs
256 (see also Section 3). The covariance matrix of measurement errors, \mathbf{B}_Y , is here considered to be
257 diagonal with non-zero terms equal to the observation error variance, σ_i^2 (Carrera and Neuman,
258 1986). Denoting by $\hat{\mathbf{Y}} = [\hat{Y}_1, \dots, \hat{Y}_N]$ the vector of model predictions at locations where data are

259 available, the ML estimate $\hat{\mathbf{X}}$ of the vector of the M uncertain model parameters can be obtained
 260 by minimizing with respect to \mathbf{X} the negative log likelihood criterion:

$$261 \quad NLL = \sum_{i=1}^N \frac{J_i}{\sigma_i^2} + \ln |\mathbf{B}_Y| + N \ln(2\pi) \quad (1)$$

262 where $J_i = (Y_i^* - \hat{Y}_i)^2$. Criterion (1) includes the weighted least square criterion (Carrera and
 263 Neuman, 1986; Bianchi Janetti et al., 2012 and references therein). Here, minimization of (1) is
 264 achieved using the iterative Levenberg-Marquardt algorithm as embedded in the well documented
 265 computational framework PEST (Doherty, 2002).

266 Alternative (competing) models which can be used to interpret available system states can be
 267 ranked by various criteria (e.g., Neuman, 2003; Ye et al., 2004, 2008; Riva et al., 2011; Bianchi
 268 Janetti et al., 2012; Ciriello et al., 2015 and references therein), including:

$$269 \quad AIC = NLL + 2M \quad (2)$$

$$270 \quad AIC_c = NLL + 2M + \frac{2M(M+1)}{N-M-1} \quad (3)$$

$$271 \quad KIC = NLL + M \ln\left(\frac{N}{2\pi}\right) - \ln |\mathbf{Q}| \quad (4)$$

272 Here, \mathbf{Q} represents the Cramer-Rao lower-bound approximation for the covariance matrix of the
 273 parameter estimates, i.e., the inverse expected Fisher information matrix, which renders a
 274 quantitative appraisal of the quality of parameter estimates and of the information content carried by
 275 data about model parameters (see, e.g., Ye et al., 2008 for details). The Akaike information
 276 criterion, AIC , is due to Akaike (1974), AIC_c to Hurvich and Tsai (1989) and KIC to Kashyap
 277 (1982). The lowest value of a given model identification criterion indicates the most favored model
 278 (according to the criterion itself) at the expense of the remaining models. Note that KIC tends to (a)
 279 penalize models proportionally to the number of their parameters, through the quantity
 280 $M \cdot \ln(N/2\pi)$ (Ye et al., 2008; Hernandez et al., 2006; Riva et al., 2011) and (b) favor models with
 281 smaller expected information content per observation, when considering models with equal

282 parameter numbers, minimum NLL values and prior probability of parameters linked to such a
 283 minimum (Ye et al., 2008). In light of these observations, we base the analyses presented in this
 284 study on KIC (4).

285 The discrimination criteria (2)-(4) can also be considered to assign posterior probability
 286 weights quantifying uncertainty associated with each of the tested models. The posterior probability
 287 linked to model M_k ($k = 1, \dots, N_M$, N_M being the number of interpreting models considered) is
 288 evaluated as (Ye et al., 2008):

$$289 \quad p(M_k | \mathbf{Y}^*) = \frac{\exp\left(-\frac{1}{2} \Delta IC_k\right) p(M_k)}{\sum_{i=1}^{N_M} \exp\left(-\frac{1}{2} \Delta IC_i\right) p(M_i)} \quad (5)$$

290 Here, $\Delta IC_k = IC_k - IC_{min}$, IC_k being either AIC (2), AIC_c (3) or KIC (4) and $IC_{min} = \min\{IC_k\}$ its
 291 minimum value calculated across the range of models examined; $p(M_k)$ is the prior probability
 292 associated with each model. One can set $p(M_k) = 1/N_M$. In case no prior information is available, all
 293 models being then characterized by the same prior probability.

294 Grounding our study on model identification criteria and the ensuing posterior probabilities (5)
 295 enables one to rank the models analyzed through their posterior probabilities and discriminate
 296 among them in a relative sense.

297 2.3.4 Multimodel Mean and Variance

298 We consider a collection \mathbf{M} of K mutually exclusive variogram models, M_k , upon which
 299 lead statistics, such as mean and variance/covariance, of NBL values are computed through Kriging
 300 at the nodes of a selected computational grid covering a given aquifer body. The models are
 301 uncertain, each of them being assigned the same prior probability $p(M_k)$. Variogram model M_k is
 302 employed in a Kriging framework to yield the mean (expectation) $E(\mathbf{Y} | \mathbf{D}, M_k)$ and the covariance
 303 $Cov(\mathbf{Y} | \mathbf{D}, M_k)$ of a vector \mathbf{Y} of random (log-transformed) NBL values, conditional on the prior
 304 data vector \mathbf{D} . The entries of the latter are evaluated according to the procedure illustrated in

305 Section 2.3.1. Averaging across the moments provided by all K variogram models renders the
 306 following (Bayesian-averaged) lead moments (Draper, 1995; Hoeting et al., 1999):

$$307 \quad E(\mathbf{Y} | \mathbf{D}) = \sum_{k=1}^K E(\mathbf{Y} | \mathbf{D}, M_k) p(M_k | \mathbf{D}) \quad (6)$$

$$308 \quad \begin{aligned} Cov(\mathbf{Y} | \mathbf{D}) &= \sum_{k=1}^K Cov(\mathbf{Y} | \mathbf{D}, M_k) p(M_k | \mathbf{D}) \\ &+ \sum_{k=1}^K [E(\mathbf{Y} | \mathbf{D}, M_k) - E(\mathbf{Y} | \mathbf{D})] \\ &\cdot [E(\mathbf{Y} | \mathbf{D}, M_k) - E(\mathbf{Y} | \mathbf{D})]^T p(M_k | \mathbf{D}) \end{aligned} \quad (7)$$

309 T denoting transpose. The conditional covariance $Cov(\mathbf{Y} | \mathbf{D})$, resulting from Bayesian model
 310 averaging (BMA), is the sum of a within- and between-model contribution. Posterior model
 311 probabilities, $p(M_k | \mathbf{D})$, are calculated according to (5) and weigh the contribution of model M_k
 312 to BMA moments. For the purpose of our demonstration, we evaluate $E(\mathbf{Y} | \mathbf{D}, M_k)$ and
 313 $Cov(\mathbf{Y} | \mathbf{D}, M_k)$ through Kriging performed upon relying on variogram model M_k characterized
 314 through ML parameter estimates.

315 **3. Results and discussion**

316 Prior to the application of the PS procedure, the correct attribution of each monitoring
 317 station to a target groundwater body has been assessed on the basis of the technical characteristics
 318 of the boreholes (e.g., position of the filters) and the available hydrogeological information (e.g., the
 319 depth of permeable layers). Note that the application of the exclusion criteria described in Section
 320 2.3.1 has been performed by disregarding NH_4 because the collected sample cores provide evidence
 321 of natural occurrence of paleo-peats (Amorosi et al., 1996; Cremonini et al., 2008) consistent with
 322 documented increasing NH_4 concentrations with depth in the three reservoirs investigated (see also
 323 Molinari et al., 2012).

324 Exclusion of samples associated with anthropogenic influence (according to the criteria
 325 listed above) is followed by identification and removal of outliers from the remaining data set,

326 yielding the temporal record subsequently employed for local NBL estimation at each control point.
 327 Outliers are here identified as high concentration values that lie outside an interval centered around
 328 the sample mean of the data remaining after pre-selection and of width equal to three times the
 329 associated standard deviation. Excluding these records from the analyses is consistent with the
 330 observation that the bulk local environmental behavior is not significantly influenced by isolated (in
 331 time) and significantly large concentrations recorded, for instance, as a result of unusual processes
 332 taking place only in a particular year (and linked to the occurrence of, e.g., extreme rains that can
 333 cause large infiltration enhancing remobilization and diffusion of natural compounds), as opposed
 334 to a typical environmental baseline observed across the entire monitoring period.

335 Application of PS to the time series of each monitoring well is inevitably linked to a
 336 reduction of the number of records to be then subject to statistical analyses. We then ground our
 337 NBL calculations solely on monitoring wells which, following data selection, exhibit a time series
 338 with more than five records, characterizing an active monitoring period spanning at least 3 years
 339 (note that samples are collected on a six-month basis). Note that reliance on this kind of analysis is
 340 warranted for aquifers where a significant number of monitoring points is available with (a) a
 341 reliable and extensive temporal record of observations and (b) an appropriate degree of spatial
 342 coverage of the system. Table 3 lists the number of observation wells effectively employed for each
 343 water body after data selection.

344

345

Table 3

346

Number of observation wells available for spatial analysis of NBLs after data selection.

groundwater body	monitoring stations	monitoring stations
	Ammonium	Arsenic
0610	51	50
0630	62	60
2700	47	45

347

348

349

When compared against Table 2, these data reveal that the upper confined water body 0610 suffers from a significant reduction (about 50%) of the total number of monitoring points due to

350 exclusion of data evidencing possible anthropogenic impact on the water body. Otherwise, only a
 351 limited reduction of observation wells is observed in the case of the other two water bodies (with a
 352 reduction of 20% and 16%, respectively for water body 0630 and 2700). With reference to
 353 anthropogenic pressures, these results are consistent with the observation that (a) the water body
 354 2700 is a deep reservoir and hence subject to more limited anthropogenic stresses than the other two
 355 bodies, which constitute upper confined reservoirs, (b) groundwater body 0630 is located in the
 356 proximity of the Po river, that can act as a source of recharge of the reservoir, thus reducing the
 357 impact of stresses caused, for instance, by pumping activities, and (c) reservoir 0610 is subject to a
 358 enhanced anthropogenic impact due to exploitation associated with intensive industrial/agricultural
 359 activities.

360 *3.1. Ammonium*

361 Values of local NBL_{90} obtained at each monitoring well through the methodology described
 362 in Section 2.3 display a high degree of spatial variability. These range from 0.828 to 20.835 mg/L
 363 (i.e., -0.08 to 1.319 in logarithmic scale) within groundwater body 0610, while ranging from 0.025
 364 to 14.406 mg/L (i.e., -1.602 to 1.159 in logarithmic scale) and from 0.025 to 30.362 mg/L (i.e., -
 365 1.602 to 1.482 in logarithmic scale), respectively within groundwater body 0630 and 2700. Values
 366 of NBL_{90}^{PS} estimated through the standard PS procedure at the regional scale (i.e., one value for each
 367 reservoir) for the three water bodies analyzed (Molinari et al., 2012) are listed in Table 4 together
 368 with the corresponding logarithmic-transformed values. We observe that all estimates exceed EU
 369 Drinking Water Standard for ammonium.

370

371 **Table 4**
 372 Values for $NH_4 NBL_{90}$ estimated via the original PS procedure (NBL_{90}^{PS}). EU Drinking Water
 373 Standard for NH_4 concentration is 0.5 mg/L.

groundwater body	NBL_{90}^{PS} (mg/L)	$\text{Log}_{10} NBL_{90}^{PS}$
0610	4.6	0.66
0630	5.2	0.71
2700	12.0	1.08

374

375 For the purpose of our application, classical isotropic Spherical or Exponential variogram
 376 models, with or without nugget, are considered. Table 5 lists the results of variogram model
 377 calibration analysis. For each groundwater body, these include the estimated set of model
 378 parameters (i.e., nugget, range, and sill) based on minimization of NLL (1) and the associated
 379 posterior probabilities calculated through (4) and (5). As stated, the analyses have been performed
 380 by considering logarithmic transformation of values of NBL_{90} (termed as $NH_4Log_{10} NBL_{90}$).

381

382

383

384 **Table 5**
 385 Estimated parameters of variogram models of $NH_4Log_{10} NBL_{90}$ values and associated posterior
 386 probability (p) based on KIC (4). Here, Sph = Spherical model, Exp = Exponential model; n =
 387 nugget, a = range (practical range for Exponential model), c = sill; (*) NBL_{90} values are given in
 (mg/L).

	Groundwater body 0610				Groundwater body 0630				Groundwater body 2700			
Model	n	a	c	p	n	a	c	p	n	a	c	p
	(*)	(km)	(*)	(%)	(*)	(km)	(*)	(%)	(*)	(km)	(*)	(%)
Sph	-	18.7	0.106	52	-	83.5	0.496	40	-	81.7	0.485	6
Exp	-	22.9	0.107	48	-	138.2	0.581	49	-	100.4	0.509	34
Sph	-	-	-	-	0.045	87.9	0.452	4.2	0.1	101.0	0.401	5
Exp	-	-	-	-	0.031	159.6	0.574	6.6	0.1	208.9	0.538	55

388

389 The spatial variogram structure of $NH_4Log_{10} NBL_{90}$ within groundwater body 0610 has been
 390 interpreted by a Spherical and an Exponential model without nugget effect (introducing a nugget
 391 effect in the model calibration procedure is associated with near-zero estimated nugget and
 392 posterior weight and is therefore disregarded). Both variogram models exhibit similar values of
 393 range (practical range in the case of Exponential model) and sill, with nearly coinciding posterior
 394 probabilities.

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Four theoretical models, i.e., Spherical and Exponential models with and without nugget, are
 considered for groundwater bodies 0630 and 2700. The contribution of the nugget to the total
 variance is less than 10% in the case of groundwater body 0630. Otherwise, the estimated nugget

398 effect for groundwater body 2700 is equal to 15.7% and 20% of the total variance, respectively for
399 the Exponential and Spherical model with nugget, suggesting the occurrence of a significant degree
400 of variability between sample pairs at short distances.

401 Estimated values of correlation scale (i.e., the range) are lowest for water body 0610, a
402 finding that might be related to the spatial arrangement of sampling boreholes across the system or
403 to the observation that state variables, such as given quantiles of concentration time series observed
404 at multiple wells, are not necessarily characterized by a strong degree of spatial correlation, due to
405 the dynamics associated with the key processes driving chemical migration in the system.
406 Otherwise, a consistent spatial persistence of correlation structure of ammonium NBL_{90} is observed
407 in water bodies 0630 and 2700, where spatial variations of kriged estimates are then expected to be
408 quite smooth. The largest posterior probability values for water body 0630 are associated with
409 Spherical and Exponential models without nugget. On the other hand, the Exponential variogram
410 model (with or without nugget) is unambiguously favored in water body 2700.

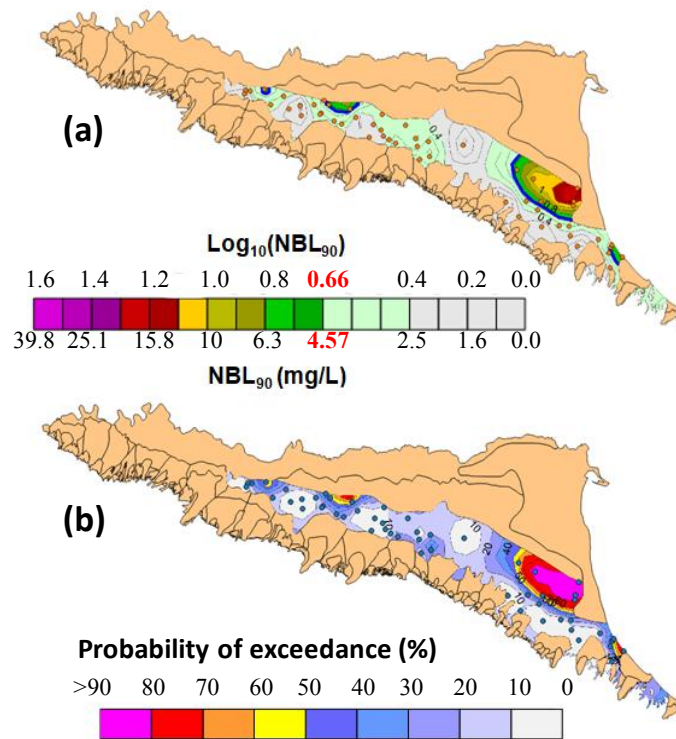
411 As described in Section 2.3, the calibrated variogram models are then employed to obtain
412 spatial distributions of Kriging estimates of log-transformed values of NBL_{90} and their associated
413 variance across each of the water bodies considered. For the purpose of this study we rely on
414 Ordinary Kriging, other flavors of Kriging being compatible with our approach. Kriging estimates
415 and variances obtained for each model are then weighted on the basis of posterior probability values
416 according to (6) and (7), to yield spatial distributions of estimated mean and variance of $\text{NH}_4\text{Log}_{10}$
417 NBL_{90} based on the complete set of tested models and fully including information on model
418 uncertainty.

419 As an example of the results obtained by considering individual variogram models, spatial
420 distributions of kriged values for local NBLs of ammonium in water body 2700 with the four
421 models listed in Table 5 are depicted in Appendix A. The spatial maps of kriged values of
422 $\text{NH}_4\text{Log}_{10}$ NBL_{90} for groundwater bodies 0610, 0630 and 2700 and resulting from our multimodel
423 analysis are respectively depicted in Figures 2, 3 and 4 together with the probability of exceeding

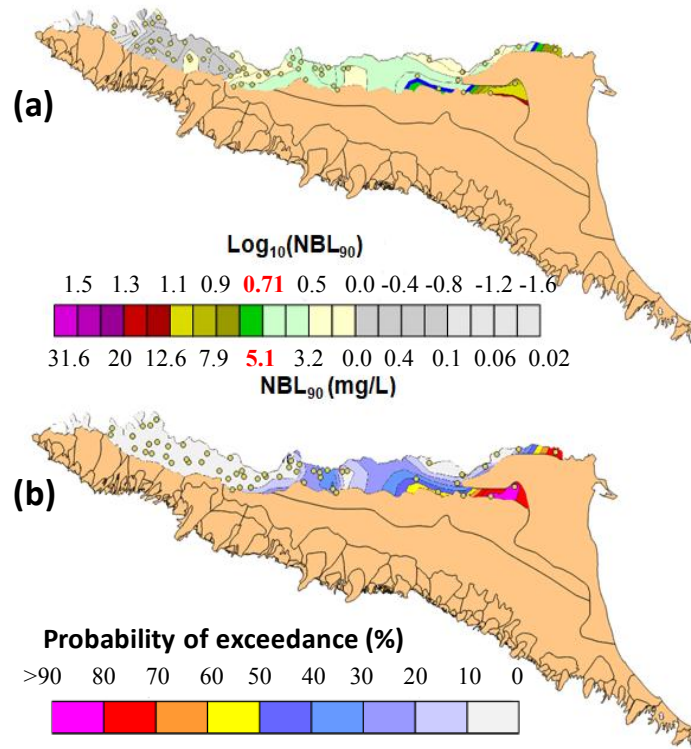
424 the uniform regional (log-transformed) NBL_{90}^{PS} value calculated via the original PS method (see
425 Table 4). For completeness, the corresponding spatial distributions of Kriging variance for the three
426 water bodies investigated are depicted in Appendix B. Exceedance probabilities are calculated by
427 assuming a Gaussian distribution of log-transformed NBL_{90} with local mean and variance rendered
428 by the multimodel analysis. Areas characterized by estimated mean values of $NH_4Log_{10}NBL_{90}$
429 larger than their uniformly distributed counterpart based on the original PS method are also
430 demarcated in Figures 2a, 3a, and 4a.

431 Spatial patterns of mean $Log_{10} NBL_{90}$ display a marked degree of variability across the
432 systems, suggesting that relying on a single (uniform) value can shadow the proper representation
433 of the actual NBL distribution within a given reservoir. For example, one can observe the
434 occurrence of areas where Kriging-based NBL_{90} exhibit values larger than NBL_{90}^{PS} . Otherwise, one
435 can also observe the occurrence of regions where estimated average NBLs are lower than NBL_{90}^{PS} ,
436 yet still larger than the EU drinking water standard set for ammonium. These observations reinforce
437 the idea that the use of a single NBL value as representative of the whole reservoir may lead to
438 misleading conclusions with reference to the natural behavior of the water body. For example, one
439 might argue that values larger than NBL_{90}^{PS} be associated with external causes (e.g., anthropogenic
440 activities) while they could be linked to specific and local hydrogeochemical natural processes.
441 Likewise, one should also consider that it is possible that concentration values in some areas be
442 lower than NBL_{90}^{PS} , yet larger than their regulatory-based counterpart (i.e., 0.5 mg/L in the case of
443 NH_4), as a result of natural processes rather than anthropogenically induced pollution phenomena.
444 These observations are also supported by the exceedance probability maps depicted in Figures 2b,
445 3b, and 4b. These suggest that there is non-negligible probability that NBL_{90}^{PS} be exceeded over a
446 large portion of the system, reinforcing the concept that the emergence of areas associated with a
447 geogenic origin of a target pollutant is masked when a single NBL_{90} value is taken as representative
448 of the whole reservoir. These sets of results represent an element upon which one could derive

449 information for the design of additional investigations. These could be directed, for example, to
 450 reduce uncertainty (as quantified, e.g., by large values of Kriging variance) in critical areas and/or
 451 to support the findings of the geostatistical analyses through the joint use of other types of
 452 information, including, e.g., (a) site-specific mineralogy, (b) the extent of petrographic provinces,
 453 or (c) the attainment of local natural hydrochemical equilibria linked to water-rock interactions with
 454 the aim of providing a complete picture of the natural signature of the system behavior. An analysis
 455 of this kind is outside the scope of the current study and will be the subject of future investigations.
 456

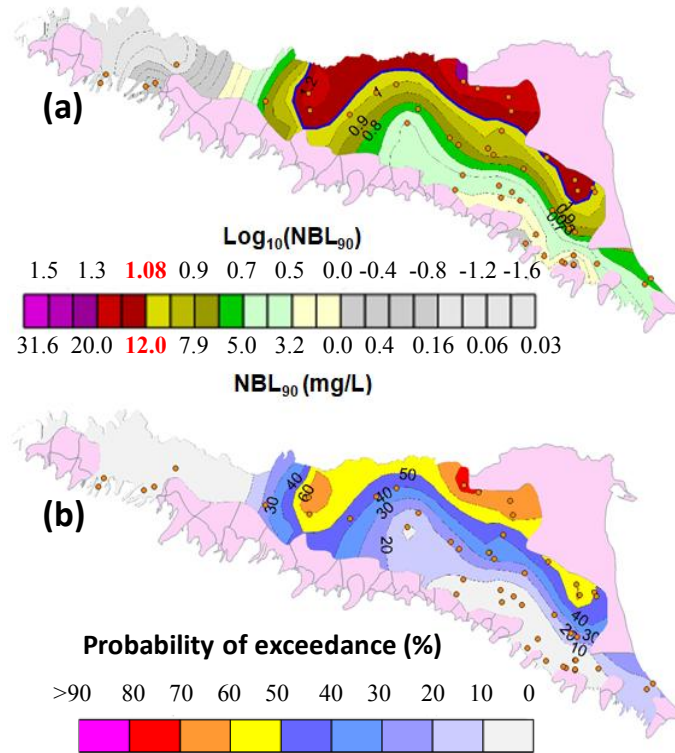


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 459 Fig. 2. Groundwater body 0610: spatial maps of (a) kriged $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}$ values and (b)
 460 probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the
 461 original PS method. The value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 0.66$ (red font on the color scale; corresponding to
 462 $\text{NBL}_{90}^{\text{PS}} = 4.6$ mg/L) is denoted by the thick blue isoline.
 463



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Fig. 3. Groundwater body 0630: spatial maps of (a) kriged $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}$ values and (b) probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the original PS method. The value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 0.71$ (red font on the color scale; corresponding to $\text{NBL}_{90}^{\text{PS}} = 5.12 \text{ mg/L}$) is denoted by the thick blue isoline.



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Fig. 4. Groundwater body 2700: spatial maps of (a) kriged $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}$ values and (b) probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the original PS method. The value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 1.08$ (red font on the color scale; corresponding to $\text{NBL}_{90}^{\text{PS}} = 12.0 \text{ mg/L}$) is denoted by the thick blue isoline.

3.2 Arsenic

481 Values of local NBL_{90} obtained for arsenic (As) through the methodology described in
482 Section 2.3 display a considerable degree of spatial variability. These range from 1 to 120.4 $\mu\text{g/L}$
483 (from 0 to 2.08 in logarithmic scale) within groundwater body 0610, and from 1 to 49.8 $\mu\text{g/L}$ (from
484 0 to 1.7 in logarithmic scale) and from 0.5 to 70 $\mu\text{g/L}$ (from -0.3 to 1.84 in logarithmic scale),
485 respectively within reservoirs 0630 and 2700.

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488
489

Table 6

Values for As NBL_{90} estimated via the original PS procedure ($\text{NBL}_{90}^{\text{PS}}$). EU Drinking Water Standard for As concentration is 10 $\mu\text{g/L}$.

groundwater body	$\text{NBL}_{90}^{\text{PS}}$ ($\mu\text{g/L}$)	$\text{Log}_{10}\text{NBL}_{90}^{\text{PS}}$
0610	33	1.52
0630	4	0.60
2700	6	0.77

490

491 Estimated values of NBL_{90}^{PS} for the three investigated water bodies (Molinari et al., 2012) are
 492 listed in Table 6 together with the corresponding log-transformed counterparts. We note that EU
 493 Drinking Water Standard for As (i.e., 10 $\mu\text{g/L}$) is exceeded only in the case of the 0610
 494 groundwater body.

495 Table 7 lists the results of the variogram model calibration analyses including, for each
 496 groundwater body, the estimated set of parameters based on minimization of NLL (1) and the
 497 associated posterior probabilities calculated through (4) and (5).

498

499

Table 7

500 Estimated parameters of variogram models of $\text{AsLog}_{10} NBL_{90}$ values and associated posterior
 501 probability (p) based on KIC (4). Here, PN = Pure Nugget, Sph = Spherical model, Exp =
 502 Exponential model; n = nugget, a = range (practical range for Exponential model), c = sill; (*)
 503 NBL_{90} values are given in ($\mu\text{g/L}$).

Model	Groundwater body 0610				Groundwater body 0630				Groundwater body 2700			
	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)
PN	0.33	-	-	100	0.252	-	-	0.002	0.238	-	-	6.7
Sph	-	-	-	-	-	12.78	0.256	0.016	-	11.57	0.243	31.6
Exp	-	-	-	-	-	14.57	0.257	0.030	-	11.57	0.243	31.6
Sph	-	-	-	-	0.132	19.80	0.125	99.77	0.07	14.42	0.172	15.0
Exp	-	-	-	-	10^{-5}	14.58	0.257	0.177	0.07	14.42	0.172	15.0

504

505 We base our results on log-transformed values of local NBL_{90} (termed as $\text{AsLog}_{10} NBL_{90}$)
 506 estimated for each control borehole. The data observed for water body 0610 are characterized by a
 507 lack of spatial correlation structure, a pure nugget effect being the only model of choice with the
 508 ability to interpret the available data. This result is consistent with the low estimated values for the
 509 range of the variogram models employed to characterize the spatial correlation structure of
 510 ammonium NBL_{90} in the same water body (see Table 5). As observed in the case of ammonium,
 511 these results might be related to (a) the lack of sampling points separated by sufficiently small
 512 length scales or (b) the nature of the variable analyzed, which might or might not display significant
 513 spatial correlation structure. As a consequence, kriged local NBL values coincide with the mean of

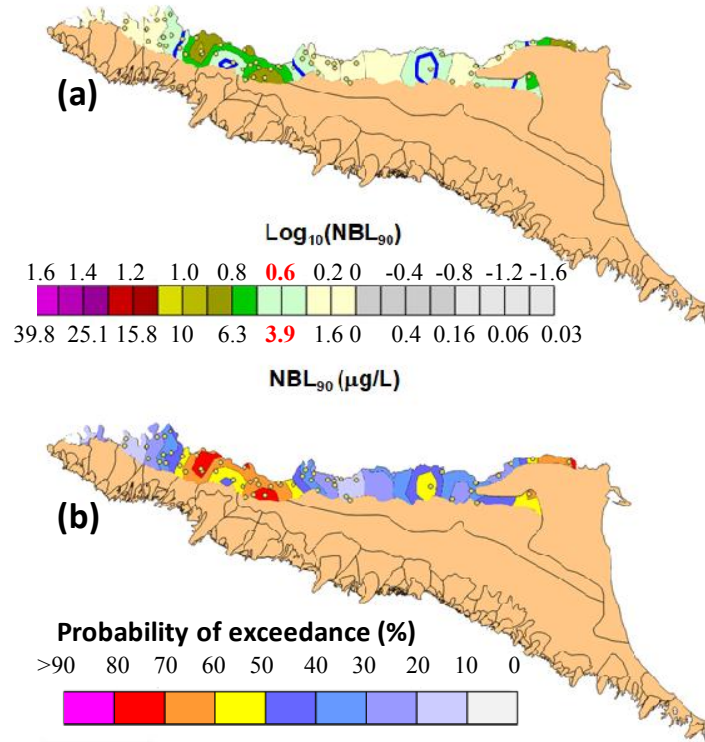
514 the available data, the associated variance being equal to the nugget. For these reasons, spatial maps
515 of $AsNBL_{90}$ for water body 0610 are not further analyzed.

516 The set of theoretical models employed to interpret the sample spatial correlation structure
517 of $AsNBL_{90}$ across groundwater bodies 0630 and 2700 comprises a pure nugget variogram model,
518 and spherical or exponential models with and without nugget. In the case of water body 0630,
519 model discrimination criteria assign the largest probabilistic weight (very close to 100%) to the
520 spherical model, which is also the model characterized by the largest contribution of the nugget to
521 the total variance (close to 50%). The other tested models are characterized by essentially negligible
522 weights. In the case of water body 2700, our results indicate that all models tested are associated
523 with a non-negligible probabilistic weight, the highest scores being equally assigned to the spherical
524 and exponential models without nugget. It is noted that addition of a nugget effect contributes to
525 about 28% of the total variance of the variogram models analyzed. Comparison of estimated
526 variogram model parameters obtained for arsenic and ammonium shows that variogram ranges for
527 As are significantly lower than those linked to NH_4 . This suggests that the spatial distribution of
528 values of $AsNBL_{90}$ may be driven by phenomena occurring at a more localized scale than in the
529 case of NH_4 .

530 We apply Ordinary Kriging to obtain Kriging estimates and variances of $AsNBL_{90}$ for each
531 variogram model. These are then weighted via posterior probability values according to (6) and (7),
532 to yield spatial distributions of estimated mean and variance of $AsLog_{10}NBL_{90}$ grounded on the
533 complete set of tested models. As an example of the results obtained by considering individual
534 variogram models, spatial distributions of kriged values for local NBLs of arsenic in water body
535 2700 with the four models listed in Table 5 are depicted in Appendix C.

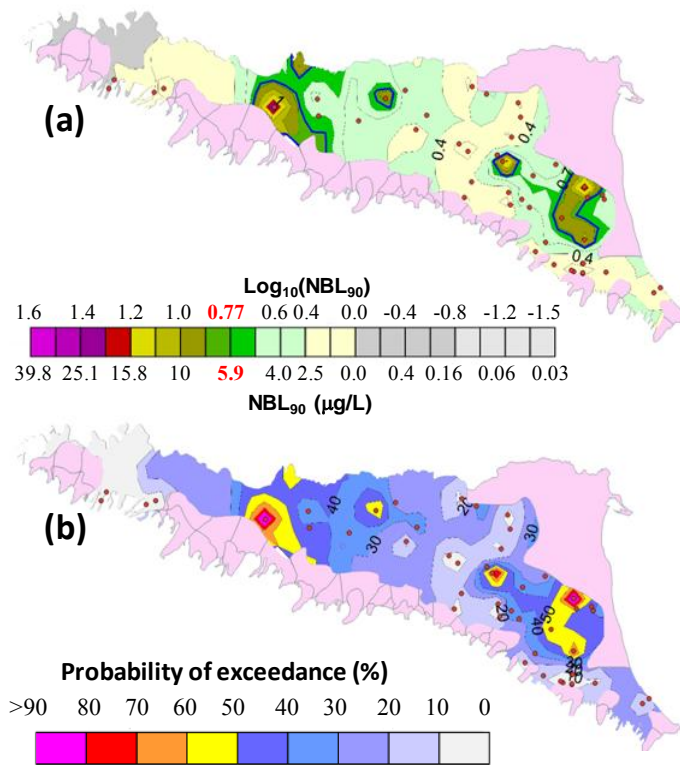
536 Kriged values of $AsLog_{10}NBL_{90}$ obtained from our multimodel analysis and associated with
537 groundwater bodies 0630 and 2700 are respectively depicted in Figures 5 and 6 together with the
538 probability of exceeding the corresponding uniform regional (log-transformed) NBL_{90}^{PS} value (see
539 Table 6). Exceedance probabilities are calculated following the procedure outlined in Section 3.1.

540 Areas characterized by estimated mean values of $AsLog_{10}NBL_{90}$ larger than their uniformly
 541 distributed counterpart based on the original PS method are also demarcated in Figures 5a and 6a.
 542 The corresponding spatial distributions of Kriging variance are depicted in Appendix D.
 543



544
 545
 546 Fig. 5. Groundwater body 0630: spatial maps of (a) kriged $AsLog_{10}NBL_{90}$ values and (b)
 547 probability of exceedance of the uniform regional (log-transformed) NBL_{90}^{PS} value calculated via the
 548 original PS method. The value $AsLog_{10}NBL_{90}^{PS} = 0.60$ (red font on the color scale; corresponding to
 549 $NBL_{90}^{PS} = 4 \mu g/L$) is denoted by the thick blue isoline.

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Fig. 6. Groundwater body 2700: spatial maps of (a) kriged $\text{AsLog}_{10}\text{NBL}_{90}$ values and (b) probability of exceedance of the uniform regional (log-transformed) $\text{NBL}_{90}^{\text{PS}}$ value calculated via the original PS method. The value $\text{AsLog}_{10}\text{NBL}_{90}^{\text{PS}} = 0.77 \mu\text{g/L}$ (red font on the color scale; corresponding to $\text{NBL}_{90}^{\text{PS}} = 6 \mu\text{g/L}$) is denoted by the thick blue isoline.

561 Similar to ammonium, we observe that considering a single NBL_{90} value as representative of
562 the whole water body can lead to underestimating/overestimating the actual natural signature which
563 can locally occur across the reservoir. These findings are markedly relevant for aquifers 0630 and
564 2700 where arsenic $\text{NBL}_{90}^{\text{PS}}$ values are lower than the EU drinking water standard, suggesting the
565 potential occurrence of localized areas characterized by local NBL values larger than the regulation
566 limit ($10 \mu\text{g/L}$). These observations are strengthened and quantified by the resulting exceedance
567 probabilities maps of arsenic (Fig. 5b and 6b). The latter could be consistent with the occurrence of
568 localized high natural arsenic content associated with the occurrence of vegetal matter, specific
569 solid fractions and redox conditions, as evidenced by Molinari et al. (2013, 2015).

569

4. Conclusions

570

Our work leads to the following major conclusions:

- 571 1. Relying on the estimate of a single (or bulk) NBL value in large-scale groundwater bodies
572 taken to represent the overall (median) behavior of the reservoir, tends to (a) mask the actual
573 distribution of local NBLs and (b) overestimate (or underestimate) natural background
574 concentrations within localized portions of the system. This might lead to misleading
575 conclusions within areas where values larger than the bulk NBL could be inappropriately
576 associated with external causes (e.g., anthropogenic activities) while they could otherwise be
577 linked to specific and localized natural processes. We show that probabilities of exceedance
578 of EU drinking water standard for hazardous species, such as arsenic, can vary quite
579 significantly across a groundwater body. These elements evidence limitations associated
580 with current directives relying on a unique NBL value taken as representative of the bulk
581 behavior of an aquifer system. Such an approach can be especially critical in large-scale
582 groundwater systems (i.e., with planar extent of the order of thousands of square kilometers)
583 of the kind we analyze in this study. Considering our strategy, which is based on the
584 application of the typical Pre-Selection (PS) methodology to the scale of each observation
585 borehole where temporal records of monitored concentrations are available, yields local
586 estimates of NBL values. These can then be embedded in a geostatistical analysis to
587 characterize their spatial variability.
- 588 2. By applying our modified PS approach, we estimate local scale NBLs, evaluated via the
589 information available at a given observation borehole. The interpretive multimodel approach
590 employed allows embedding explicitly uncertainties linked to the choice of the theoretical
591 model selected to describe the degree of spatial correlation (as embedded in the variogram
592 model) of such local scale NBL values. Relying on a single interpretive variogram model
593 can lead to an incomplete quantification of uncertainty, thus suggesting the need for
594 considering the contributions of a set of candidate models to underpin the incomplete
595 knowledge of the system behavior.

- 596 3. Spatial estimates of local NBLs and the ensuing uncertainty can serve as a basis to assist (i)
597 the demarcation of point and diffuse sources of potentially contaminating areas defining key
598 starting points and clean-up goals for remediation actions; and (ii) the management of
599 monitoring networks to optimize information content from areas associated with high
600 uncertainty or high probability of exceedance of environmental limits/standards.
- 601 4. Future European and national regulations related to NBLs assessment in large scale water
602 bodies should foresee the possibility of adopting spatial maps grounded on local NBLs
603 (based on temporal data sets). Otherwise, a unique value of NBL could be considered as
604 representative of the whole system for small scale reservoirs (planar extension lower than
605 hundreds of square kilometers) where limited amount of monitoring stations is typically
606 available and hampers a robust assessment of NBL spatial maps (with the associated
607 uncertainty).

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616 developed by the Water Challenges for a Changing World Joint Programme Initiative (Water JPI).
617 Data for the geostatistical analyses are available at the following link:
618 <https://data.mendeley.com/datasets/km6fgx227j/draft?a=3bdb20a5-adc7-429e-b1d0-3a12bd4202d8>.

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695 [circle/bridge/library?l=/deliverables/d22_final_reppdf/_EN_1.0_&a=d](http://nfp-at.eionet.europa.eu/Public/irc/eionet-circle/bridge/library?l=/deliverables/d22_final_reppdf/_EN_1.0_&a=d).

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775

Appendix A. Supplementary data

776 Spatial distributions of kriged values for local NBLs of ammonium in water body 2700 with the
777 four models listed in Table 5 (the value $\text{NH}_4\text{Log}_{10}\text{NBL}_{90}^{\text{PS}} = 1.08$ (red font on the color scale;
778 corresponding to $\text{NBL}_{90}^{\text{PS}} = 12.0$ mg/L) is denoted by the thick blue isoline.

779

780

Appendix B. Supplementary data

781 Spatial distributions of Kriging variance obtained for ammonium for water bodies 0610, 0630 and
782 2700 as a result of the multimodel analysis.

783

784

Appendix C. Supplementary data

785 Spatial distributions of kriged values for local NBLs of arsenic in water body 2700 with the four
786 models listed in Table 5. The value $\text{AsLog}_{10}\text{NBL}_{90}^{\text{PS}} = 0.77$ $\mu\text{g/L}$ (red font on the color scale;
787 corresponding to $\text{NBL}_{90}^{\text{PS}} = 6$ $\mu\text{g/L}$) is denoted by the thick blue isoline.

788

789

Appendix D. Supplementary data

790 Spatial distributions of Kriging variance obtained for arsenic for water bodies 0630 and 2700 as a
791 result of the multimodel analysis.

792

Table 1

Extension and main characteristic length scales of the three groundwater bodies investigated.

Groundwater body	Average thickness (m)	Average depth (m)	Areal extent (km ²)
0610	130	75	2928
0630	110	65	1995
2700	180	200	6934

Table 2

Number of monitoring stations and total number of samples available

Groundwater body	monitoring stations	number of samples	
		As	NH ₄
0610	90	1968	2230
0630	75	1692	1917
2700	55	1201	1383

Table 3

Number of observation wells available for spatial analysis of NBLs after data selection.

groundwater body	monitoring stations	monitoring stations
	Ammonium	Arsenic
0610	51	50
0630	62	60
2700	47	45

Table 4

Values for NH₄ NBL₉₀ estimated via the original PS procedure (NBL₉₀^{PS}). EU Drinking Water

Standard for NH₄ concentration is 0.5 mg/L.

groundwater body	NBL ₉₀ ^{PS} (mg/L)	Log ₁₀ NBL ₉₀ ^{PS}
0610	4.6	0.66
0630	5.2	0.71
2700	12.0	1.08

Table 5

Estimated parameters of variogram models of $\text{NH}_4\text{Log}_{10} \text{NBL}_{90}$ values and associated posterior probability (p) based on KIC (4). Here, Sph = Spherical model, Exp = Exponential model; n = nugget, a = range (practical range for Exponential model), c = sill; (*) NBL_{90} values are given in (mg/L).

Model	Groundwater body 0610				Groundwater body 0630				Groundwater body 2700			
	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)
Sph	-	18.7	0.106	52	-	83.5	0.496	40	-	81.7	0.485	6
Exp	-	22.9	0.107	48	-	138.2	0.581	49	-	100.4	0.509	34
Sph	-	-	-	-	0.045	87.9	0.452	4.2	0.1	101.0	0.401	5
Exp	-	-	-	-	0.031	159.6	0.574	6.6	0.1	208.9	0.538	55

Table 6

Values for As NBL_{90} estimated via the original PS procedure (NBL_{90}^{PS}). EU Drinking Water

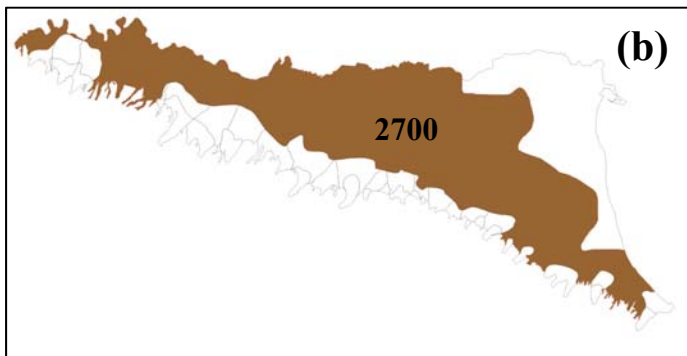
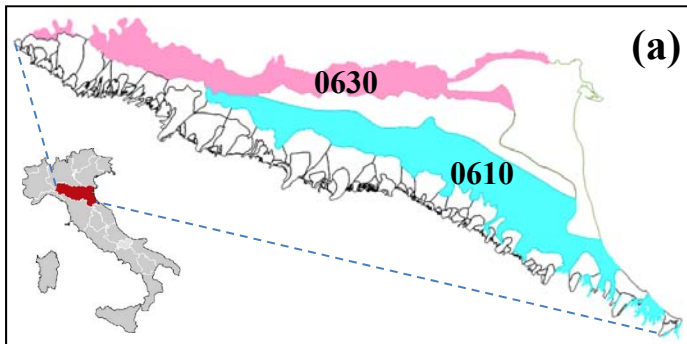
Standard for As concentration is 10 $\mu\text{g/L}$.

groundwater body	NBL_{90}^{PS} ($\mu\text{g/L}$)	$\text{Log}_{10} NBL_{90}^{PS}$
0610	33	1.52
0630	4	0.60
2700	6	0.77

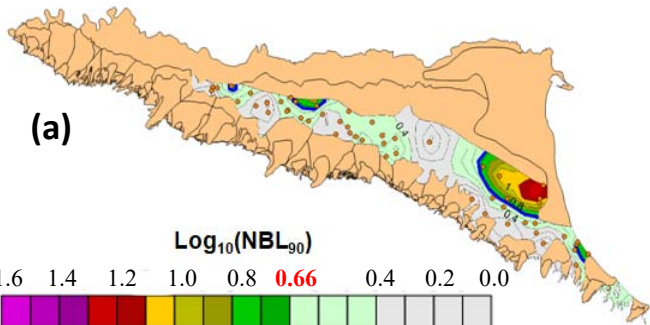
Table 7

Estimated parameters of variogram models of AsLog_{10} NBL_{90} values and associated posterior probability (p) based on KIC (4). Here, PN = Pure Nugget, Sph = Spherical model, Exp = Exponential model; n = nugget, a = range (practical range for Exponential model), c = sill; (*) NBL_{90} values are given in ($\mu\text{g/L}$).

Model	Groundwater body 0610				Groundwater body 0630				Groundwater body 2700			
	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)	n (*)	a (km)	c (*)	p (%)
PN	0.33	-	-	100	0.252	-	-	0.002	0.238	-	-	6.7
Sph	-	-	-	-	-	12.78	0.256	0.016	-	11.57	0.243	31.6
Exp	-	-	-	-	-	14.57	0.257	0.030	-	11.57	0.243	31.6
Sph	-	-	-	-	0.132	19.80	0.125	99.77	0.07	14.42	0.172	15.0
Exp	-	-	-	-	10^{-5}	14.58	0.257	0.177	0.07	14.42	0.172	15.0



(a)



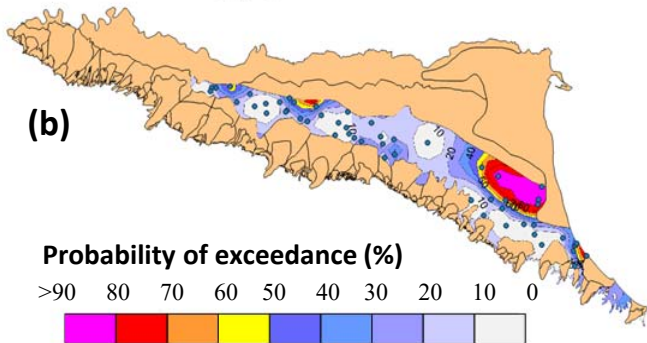
$\text{Log}_{10}(\text{NBL}_{90})$

1.6 1.4 1.2 1.0 0.8 0.66 0.4 0.2 0.0

39.8 25.1 15.8 10 6.3 4.57 2.5 1.6 0.0

NBL_{90} (mg/L)

(b)



Probability of exceedance (%)

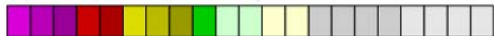
>90 80 70 60 50 40 30 20 10 0



(a)

$\text{Log}_{10}(\text{NBL}_{90})$

1.5 1.3 1.1 0.9 **0.71** 0.5 0.0 -0.4 -0.8 -1.2 -1.6



31.6 20 12.6 7.9 **5.1** 3.2 0.0 0.4 0.1 0.06 0.02

NBL_{90} (mg/L)

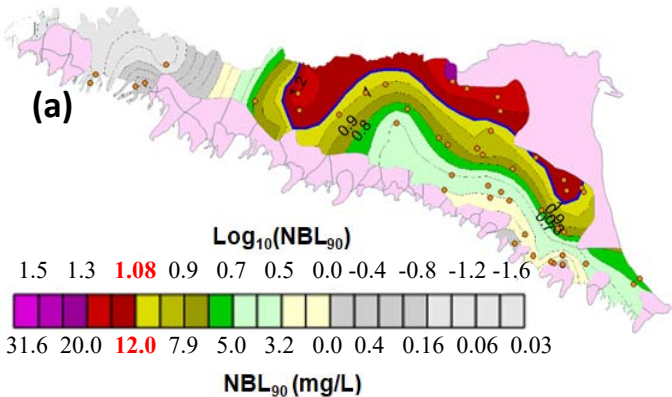
(b)

Probability of exceedance (%)

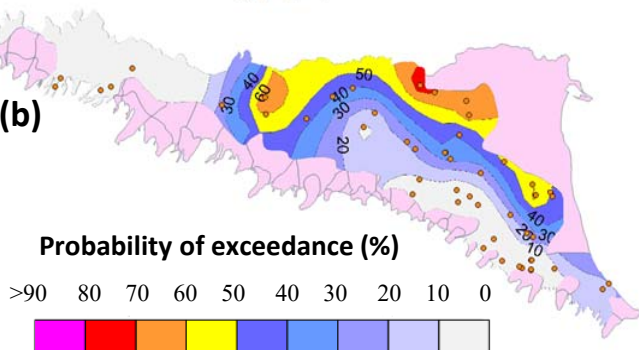
>90 80 70 60 50 40 30 20 10 0



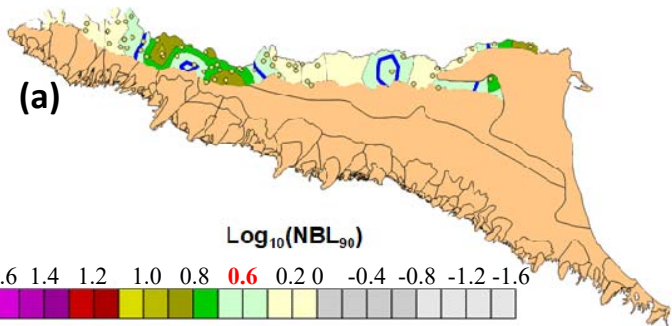
(a)



(b)

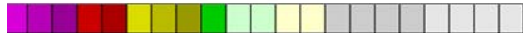


(a)



$\text{Log}_{10}(\text{NBL}_{90})$

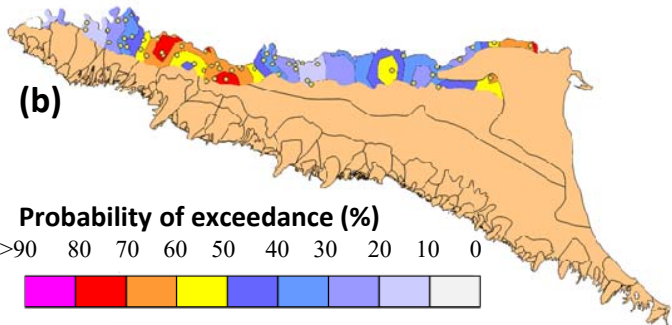
1.6 1.4 1.2 1.0 0.8 0.6 0.2 0 -0.4 -0.8 -1.2 -1.6



39.8 25.1 15.8 10 6.3 3.9 1.6 0.4 0.16 0.06 0.03

NBL_{90} ($\mu\text{g/L}$)

(b)

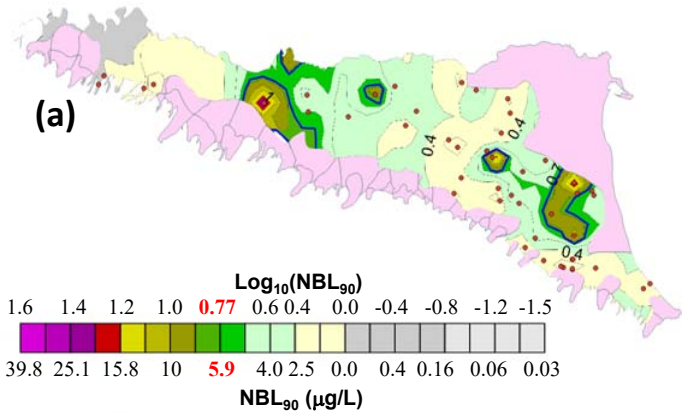


Probability of exceedance (%)

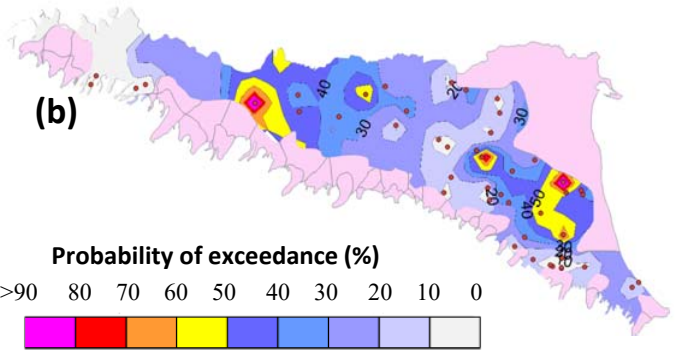
>90 80 70 60 50 40 30 20 10 0



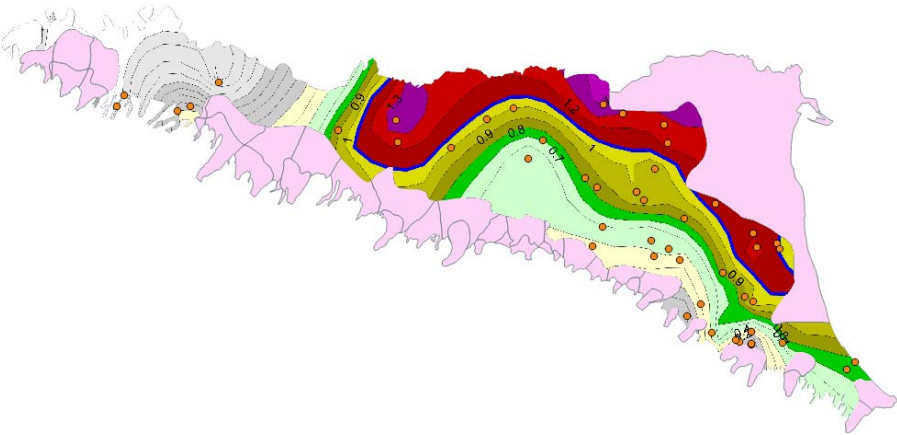
(a)



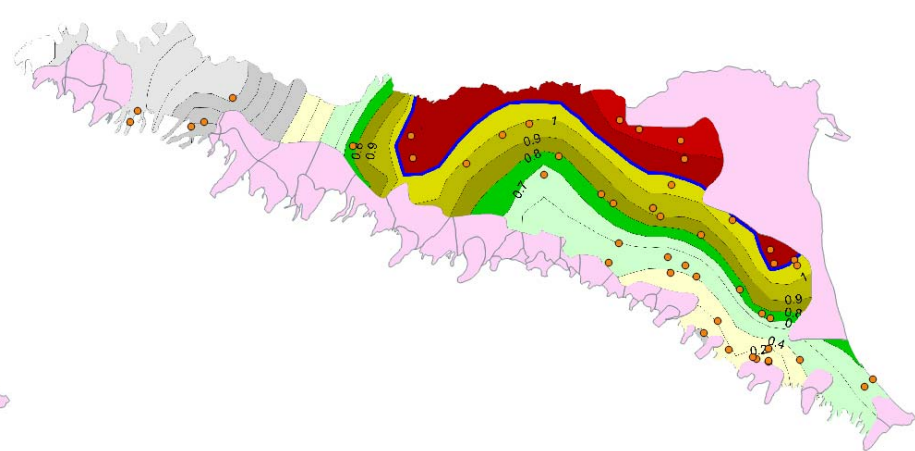
(b)



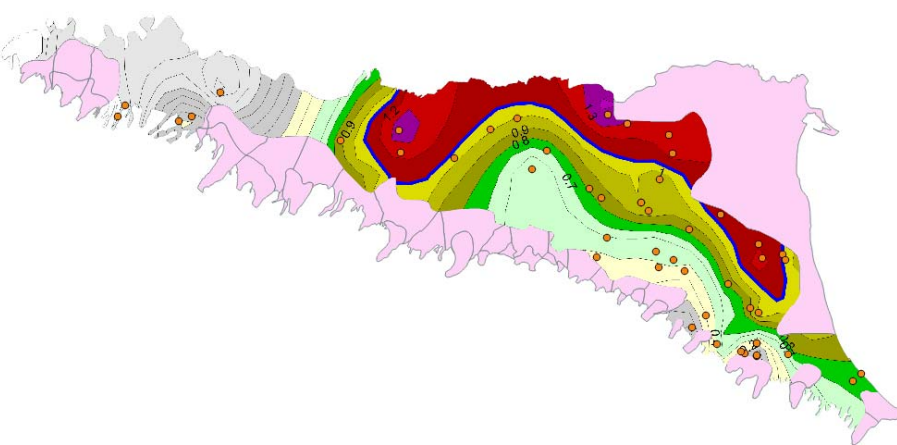
a) Spherical model without nugget



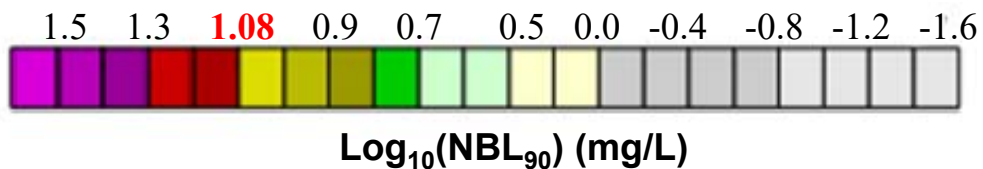
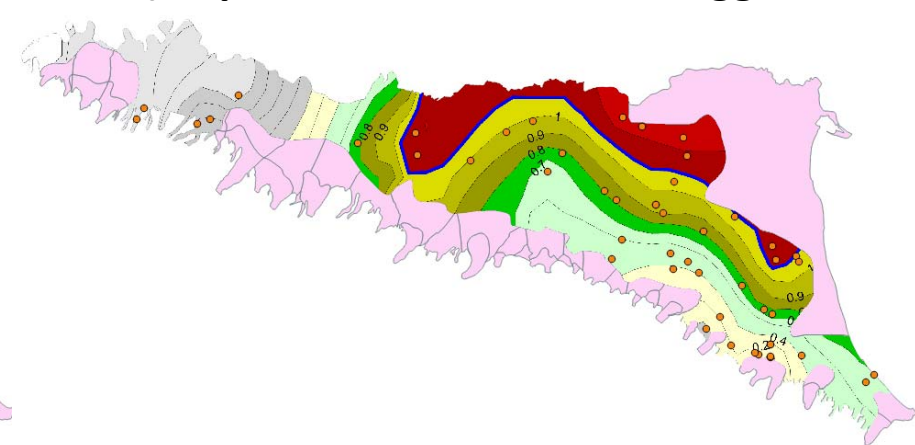
b) Spherical model with nugget

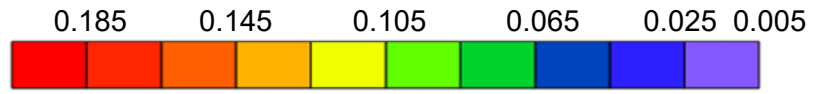


c) Exponential model without nugget



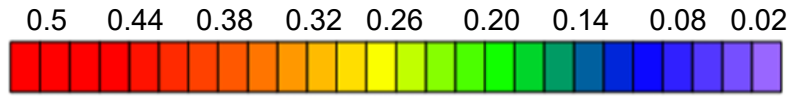
d) Exponential model with nugget





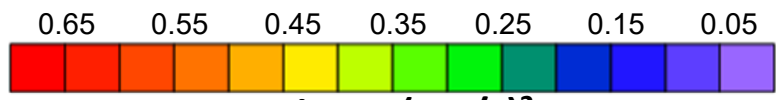
Variance (mg/L)²

Water body 0610



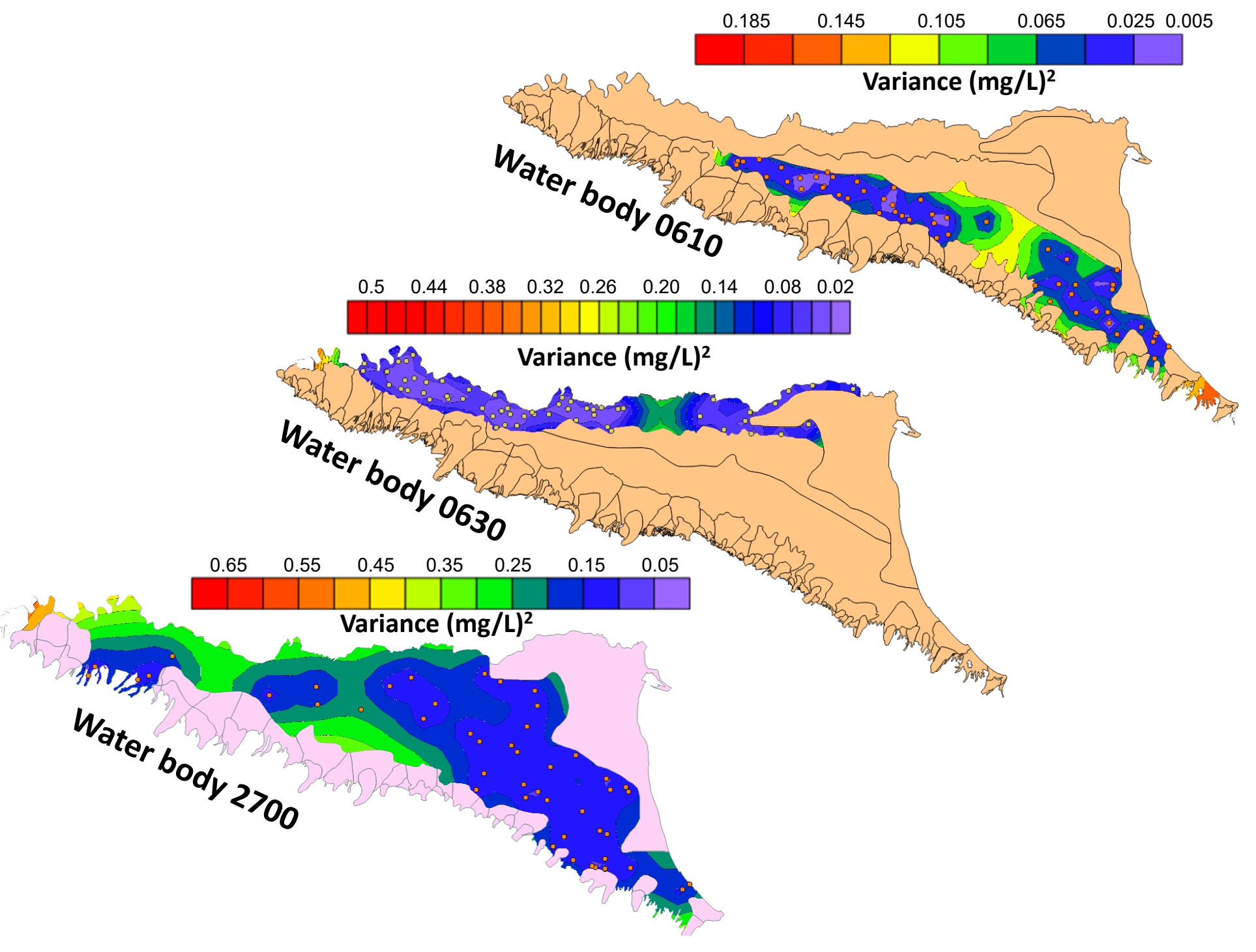
Variance (mg/L)²

Water body 0630

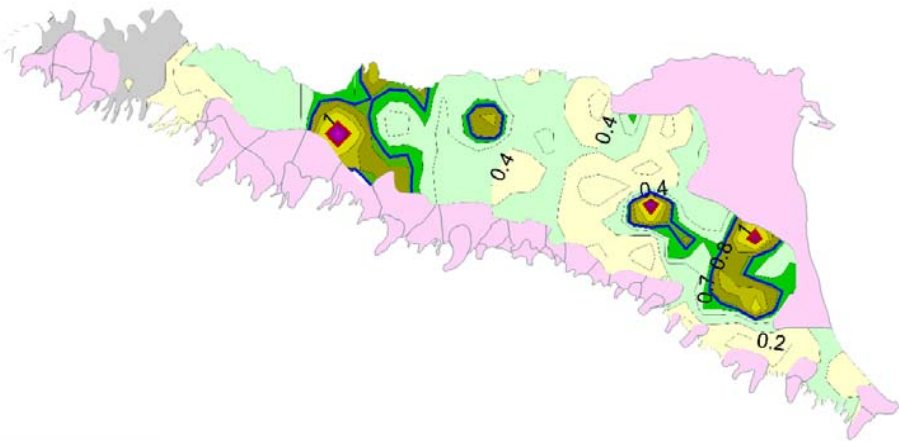


Variance (mg/L)²

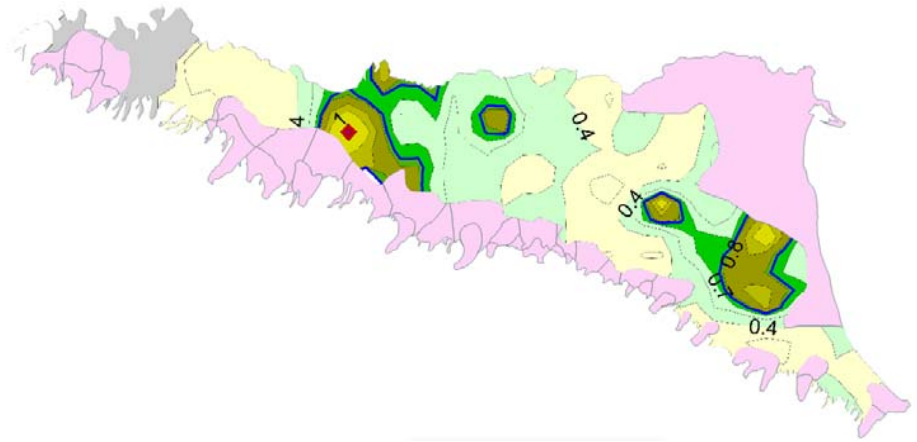
Water body 2700



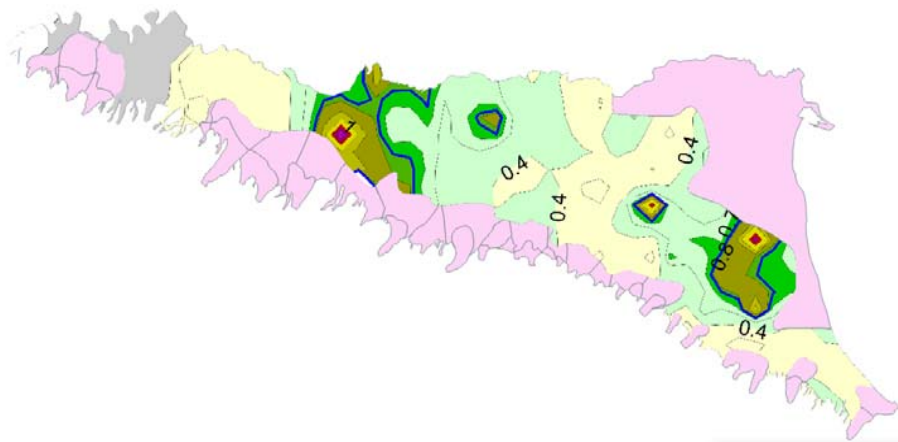
a) Spherical model without nugget



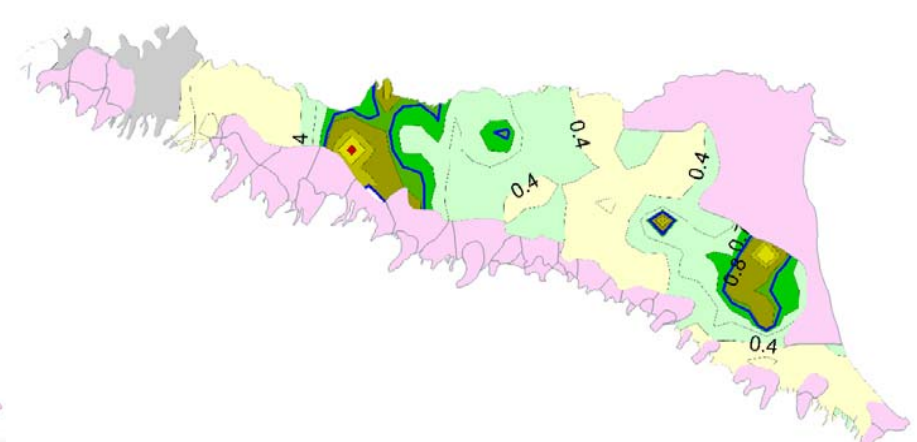
b) Spherical model with nugget



c) Exponential model without nugget

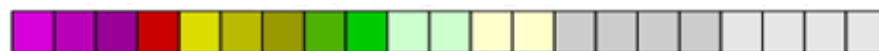


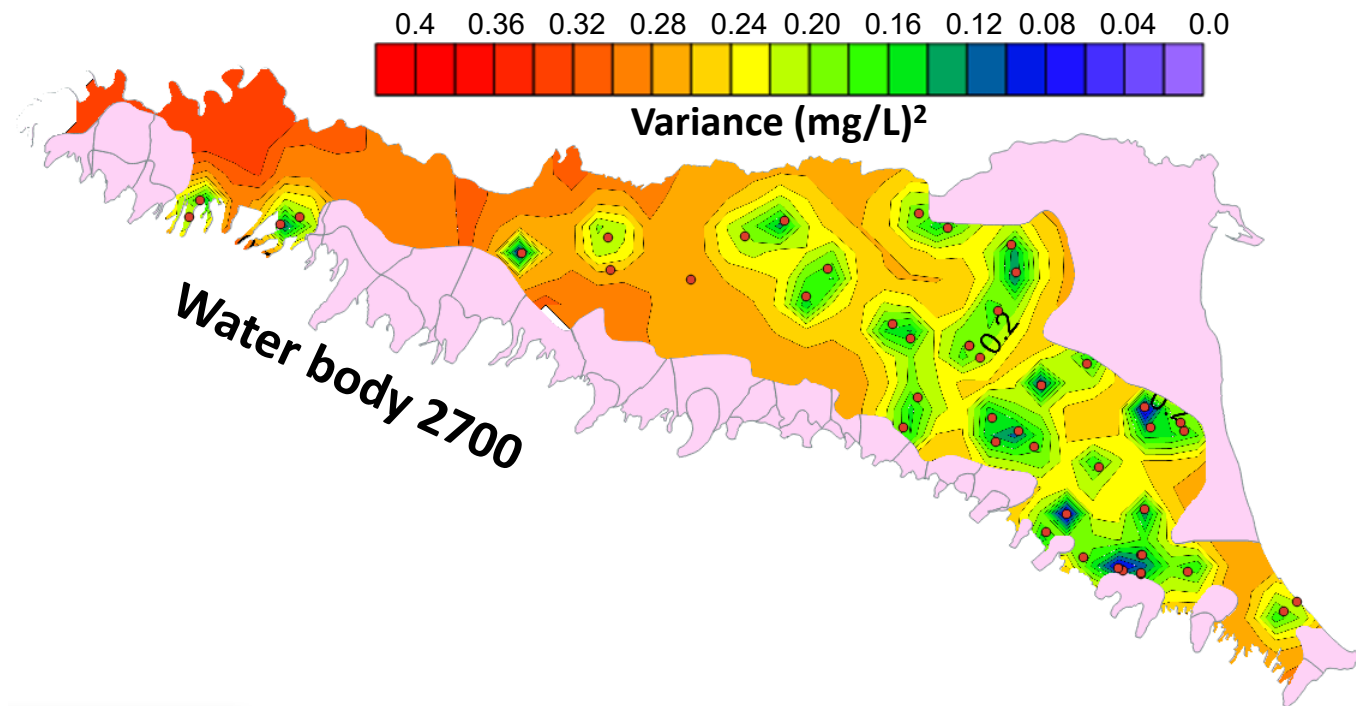
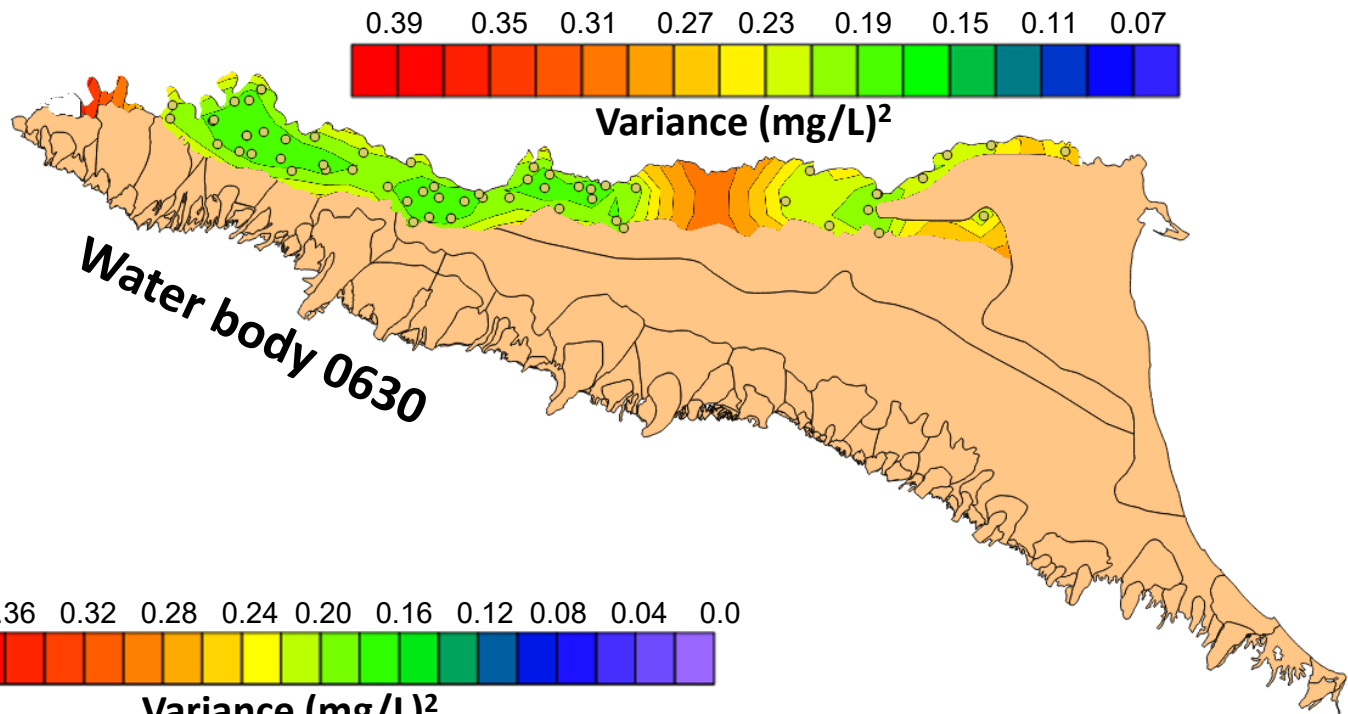
d) Exponential model with nugget



$\text{Log}_{10}(\text{NBL}_{90})$ (µg/L)

1.6 1.4 1.2 1.0 **0.77** 0.6 0.4 0.0 -0.4 -0.8 -1.2 -1.5





Electronic Supplementary Material (for online publication only)

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